

Introduction to Satellite Waves Remote sensing

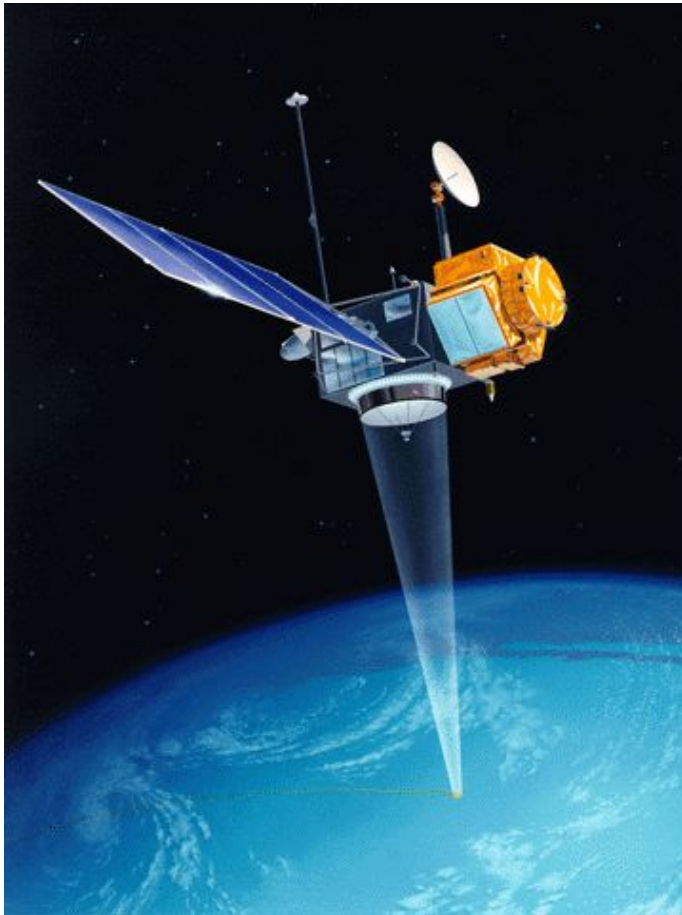
Fabrice COLLARD (OceanDataLab)

With inputs from

Jean Tournadre (IFREMER)

Fabrice Ardhuin (IFREMER)

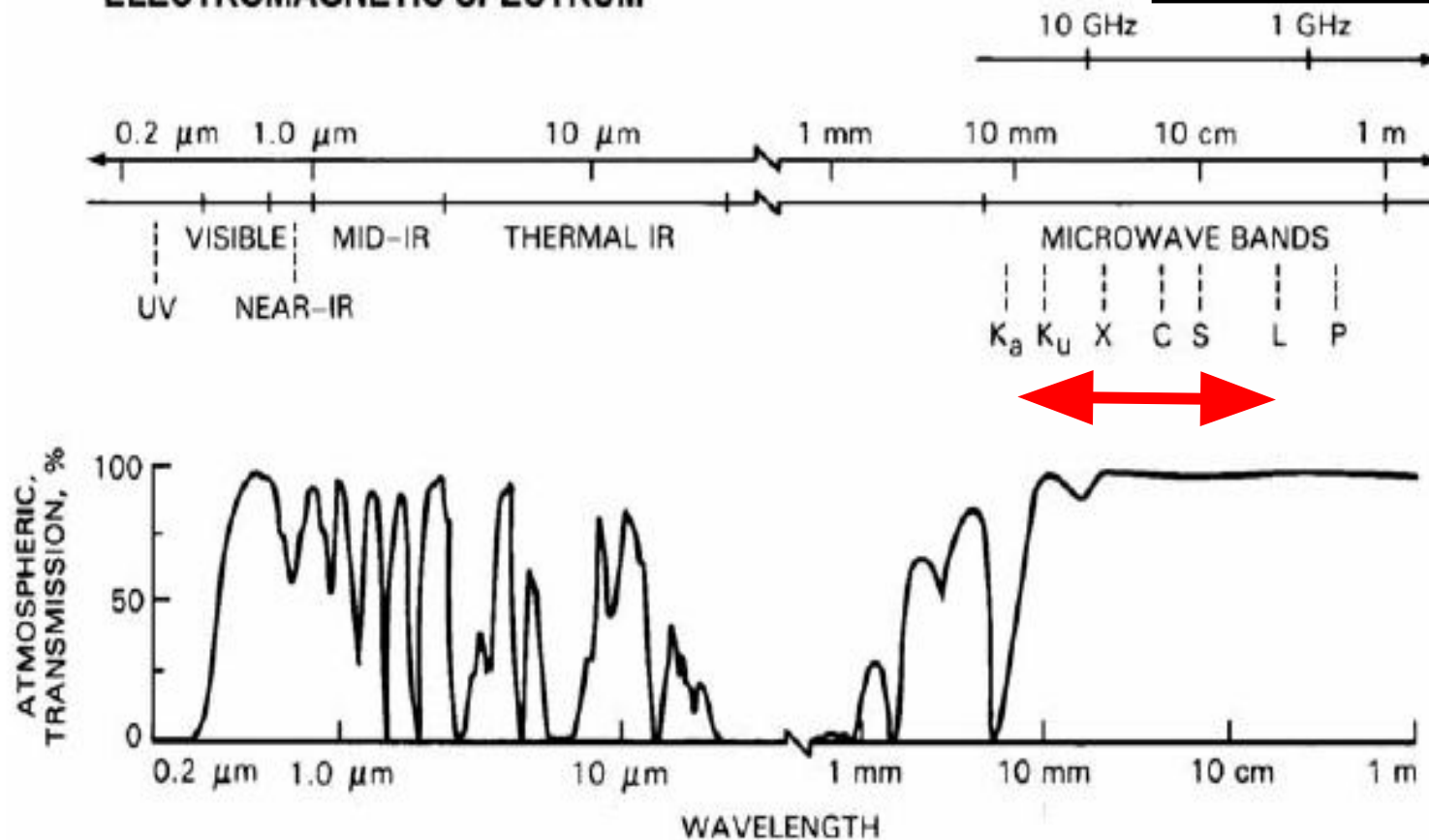
Satellite altimetry

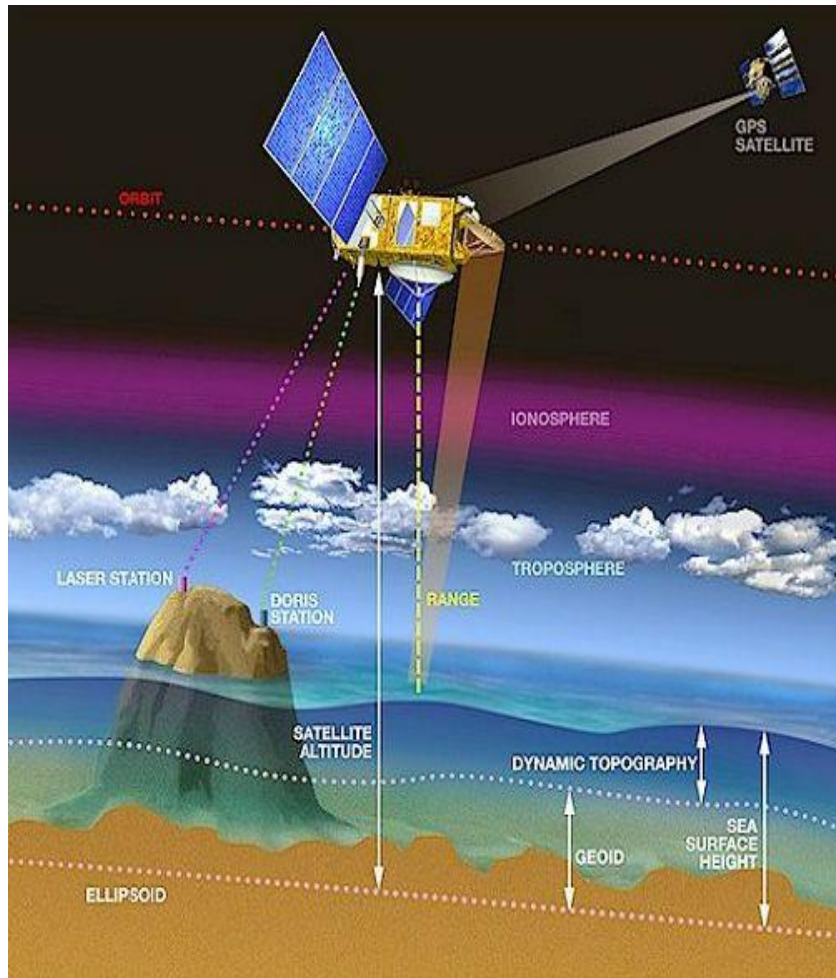


Satellite altimeters are radars, which transmit short pulses toward the earth beneath them. The return time of the pulse after reflection at the earth's surface is measured, and this yields the height of the satellite. The most important are ERS-1, ERS-2, TOPEX/Poseidon, Jason-series and Sentinel-3

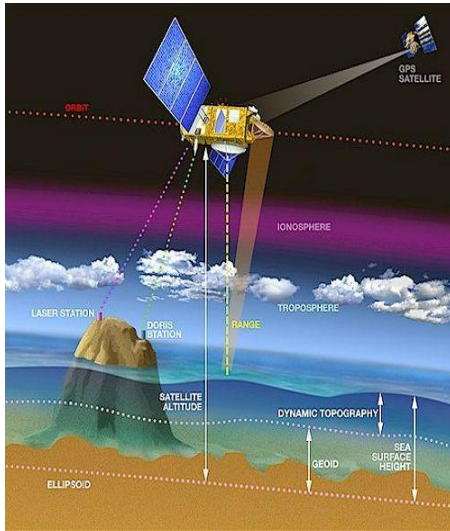
$$\lambda = c * T = c * 1/f$$

ELECTROMAGNETIC SPECTRUM



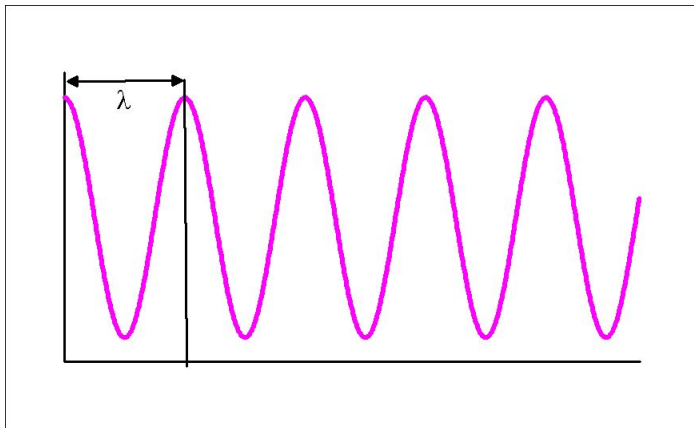


Radar altimeters on board the satellite permanently transmit signals at high frequency (Topex-Poseidon - over 1700 pulses per second) to Earth, and receive the echo from the sea surface. This is analyzed to derive a precise measurement of the round-trip time between the satellite and the sea surface. The time measurement, scaled by the speed of light (at which electromagnetic waves travel), yields a measurement of the satellite-to-ocean range.

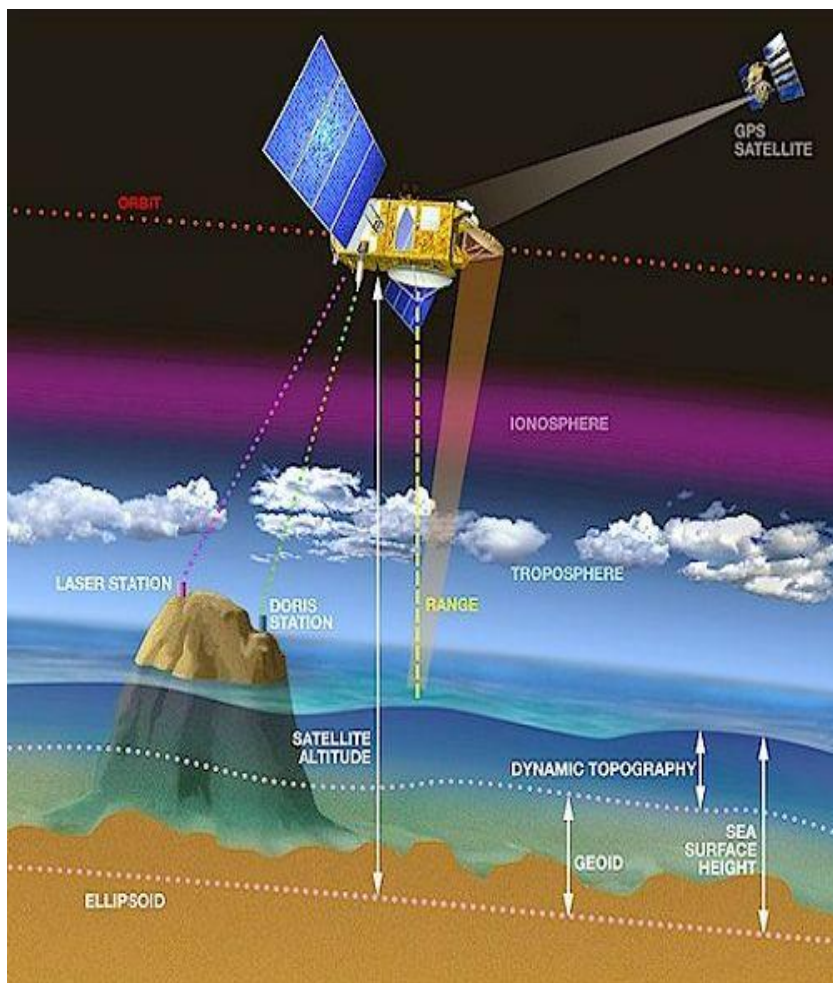


Each signal pulse produced by a radar altimeter consists of coherent electromagnetic energy.

Receiving the echo signal, the sensor measures not only the amplitude but also the phase of the reflected signal and the travel time of the pulse.



Measuring the phase of the signal reflected from the earth surface enables measuring the travel time with very high accuracy.

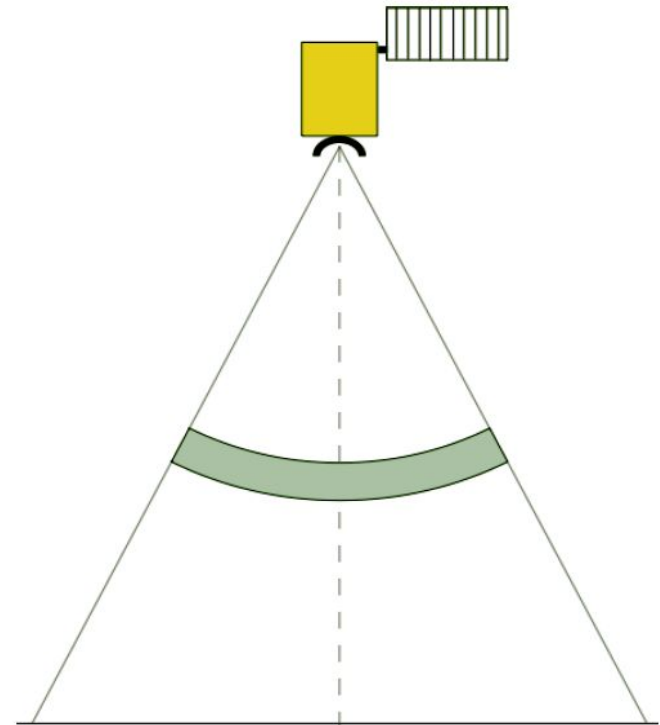
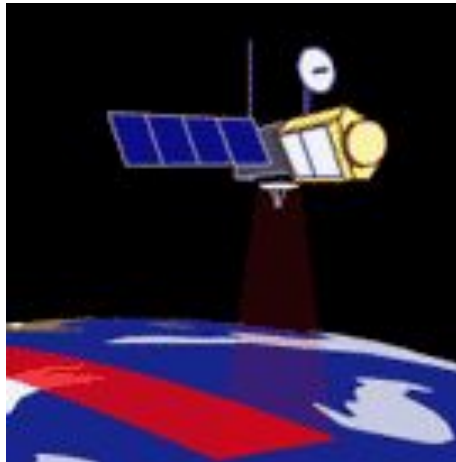


As electromagnetic waves travel through the atmosphere, they can be decelerated by water vapor or by ionization. Once these phenomena are corrected for, the final range R is estimated within 2 cm.

The ultimate aim is to measure the sea level. This requires independent measurements of the satellite orbital trajectory, i.e. exact latitude, longitude and altitude coordinates.

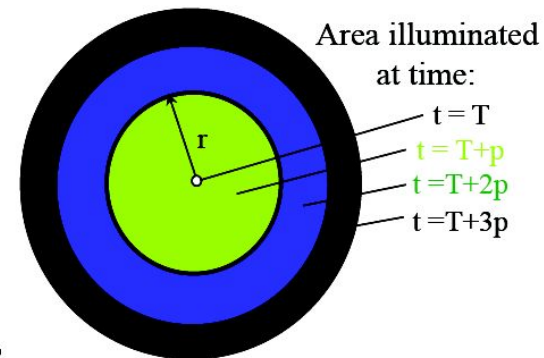
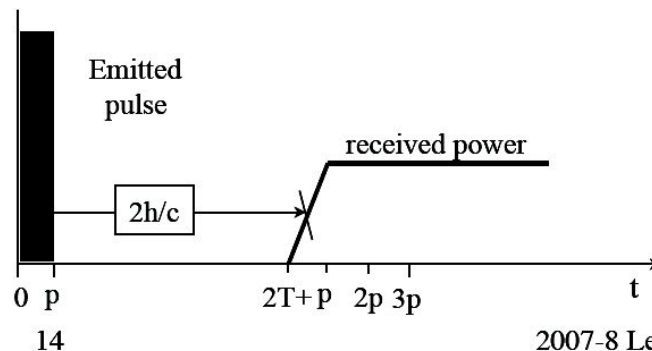
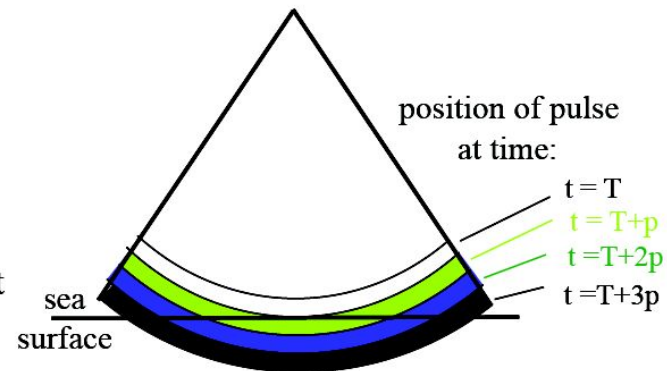
Pulse Limited Altimeter

- In a pulse limited altimeter the shape of the return is dictated by the length (width) of the pulse

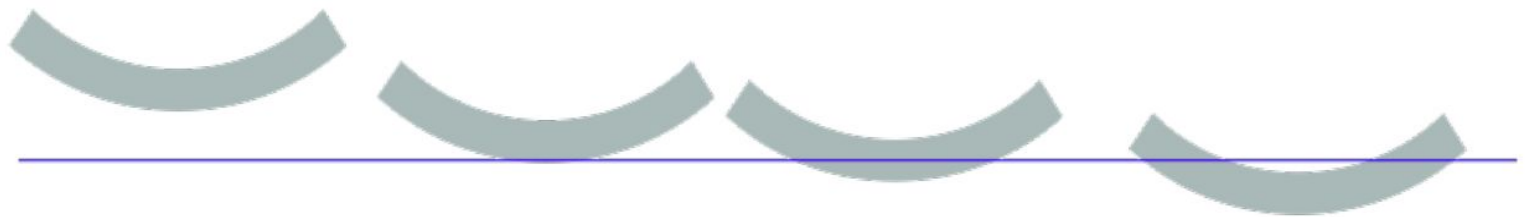


The “pulse-limited” footprint

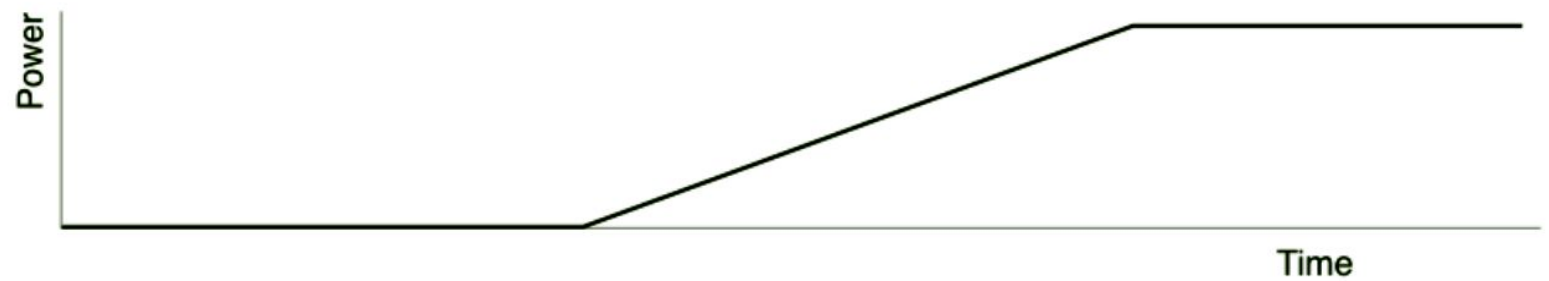
- Full illumination when rear of pulse reaches the sea – then area illuminated stays constant
- Area illuminated has radius $r = \sqrt{2hcp}$
- Measure interval between mid-pulse emission and time to reach half full height



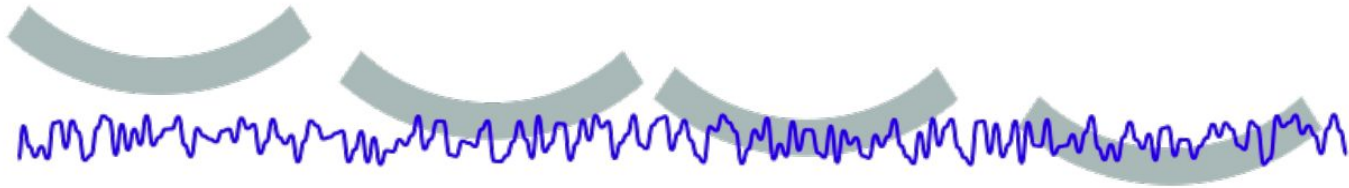
2007-8 Lecture 12



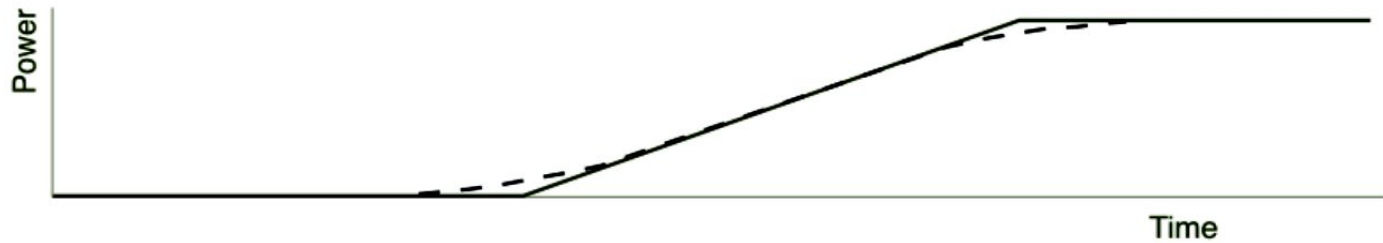
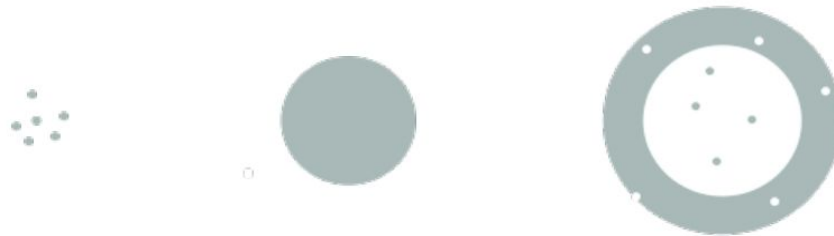
Sea Surface

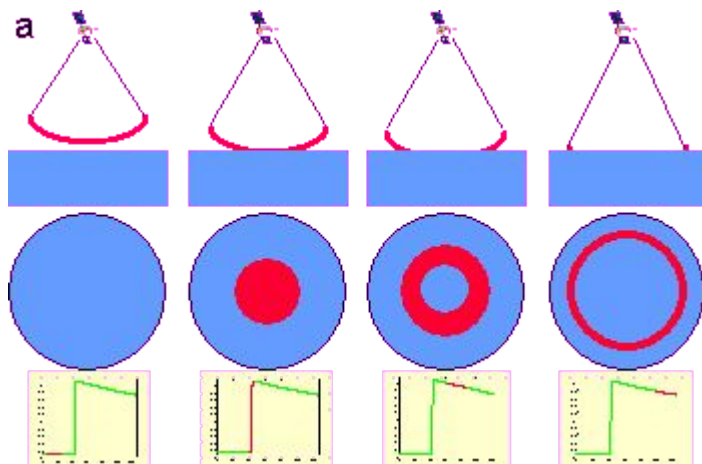


If we add waves ...

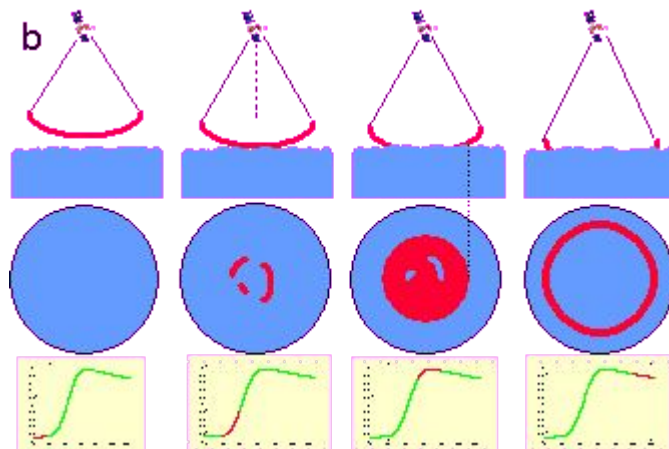


Sea Surface



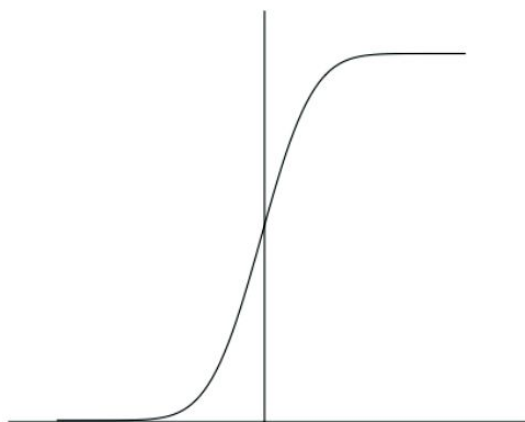


Flat surface



Wavy surface

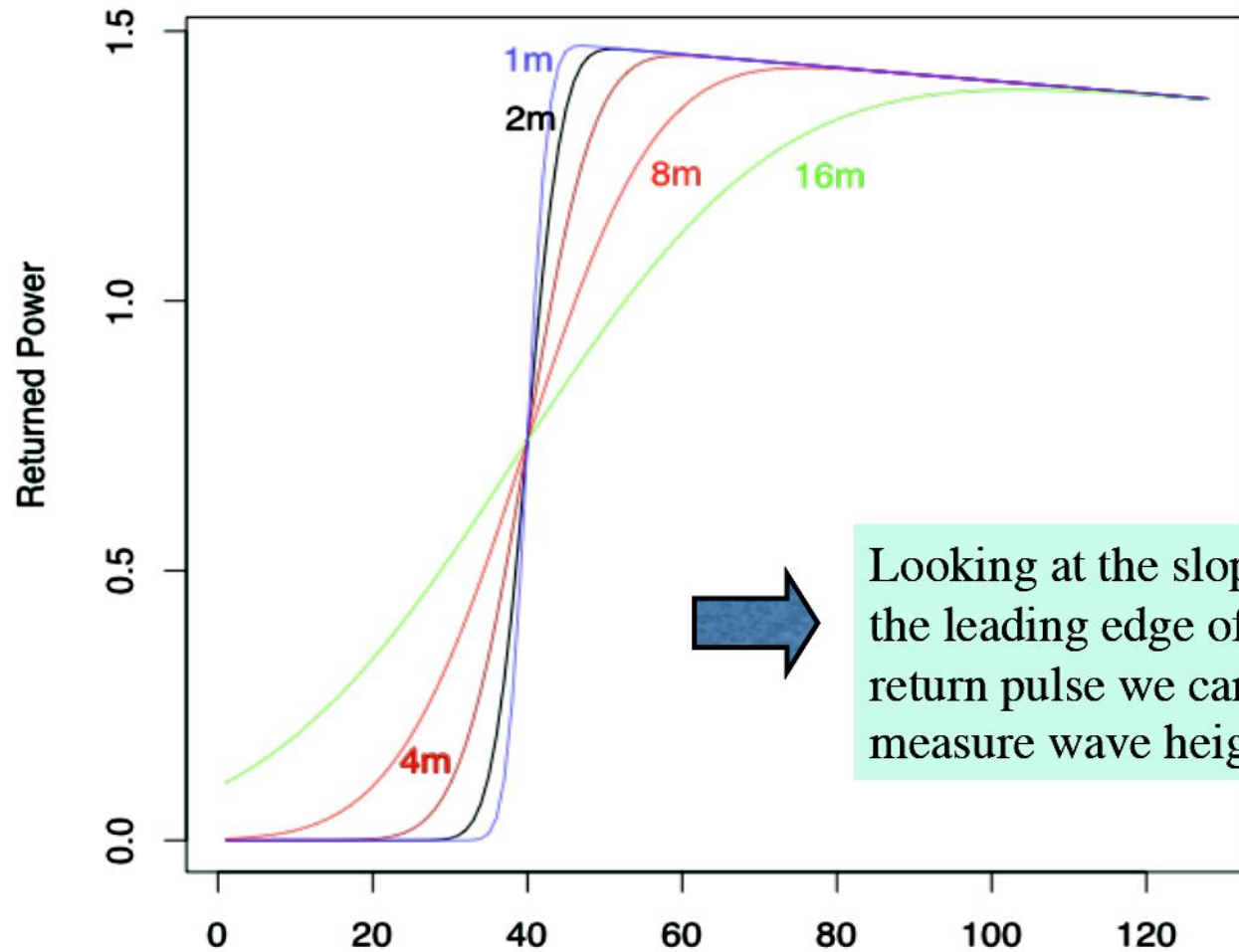
- A plot of return power versus time for a pulse limited altimeter looks like the integral of the heights of the specular points, i.e. the cumulative distribution function (cdf) of the specular scatterers



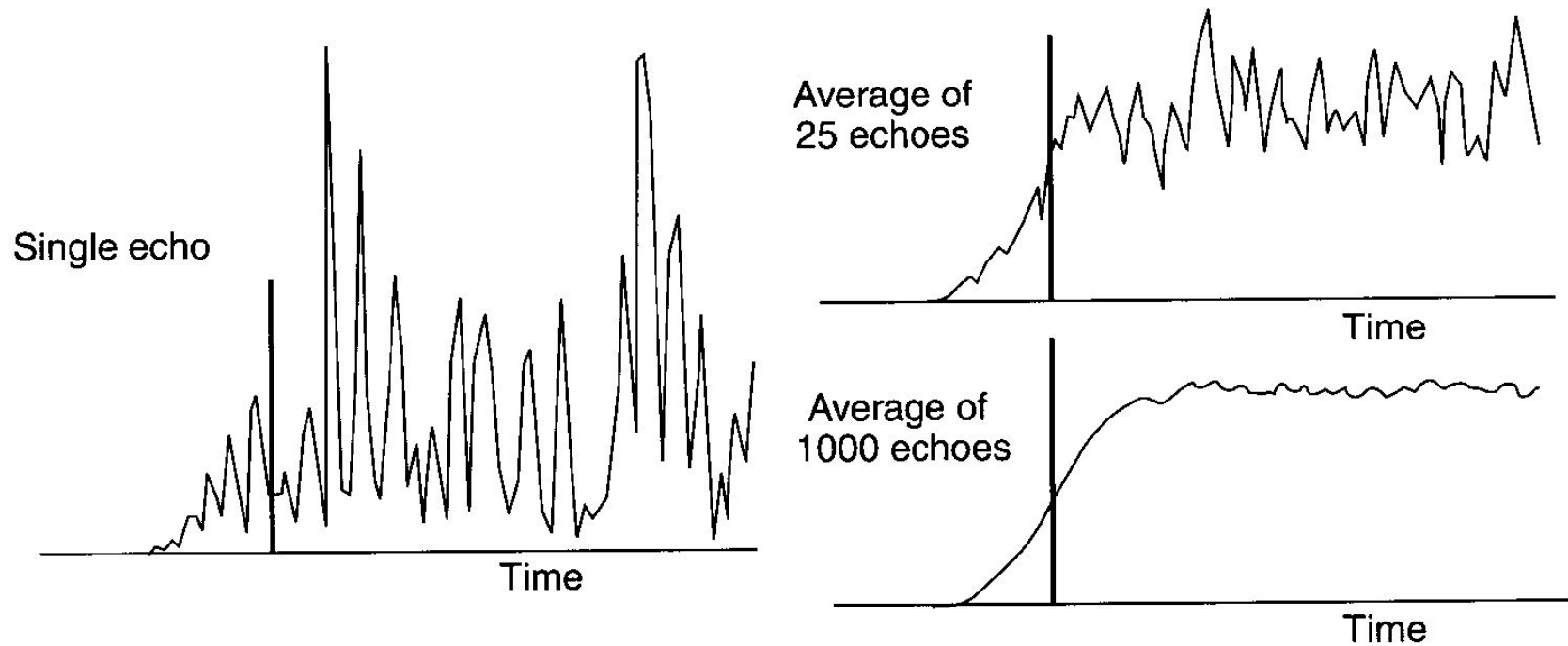
The tracking point is the half power point of the curve

Hs (m)	Effective footprint (km) (800 km altitude)	Effective footprint (km) (1335 km altitude)
0	1.6	2.0
1	2.9	3.6
3	4.4	5.5
5	5.6	6.9
10	7.7	9.6
15	9.4	11.7
20	10.8	13.4

Some example waveforms



Looking at the slope of the leading edge of the return pulse we can measure wave height!

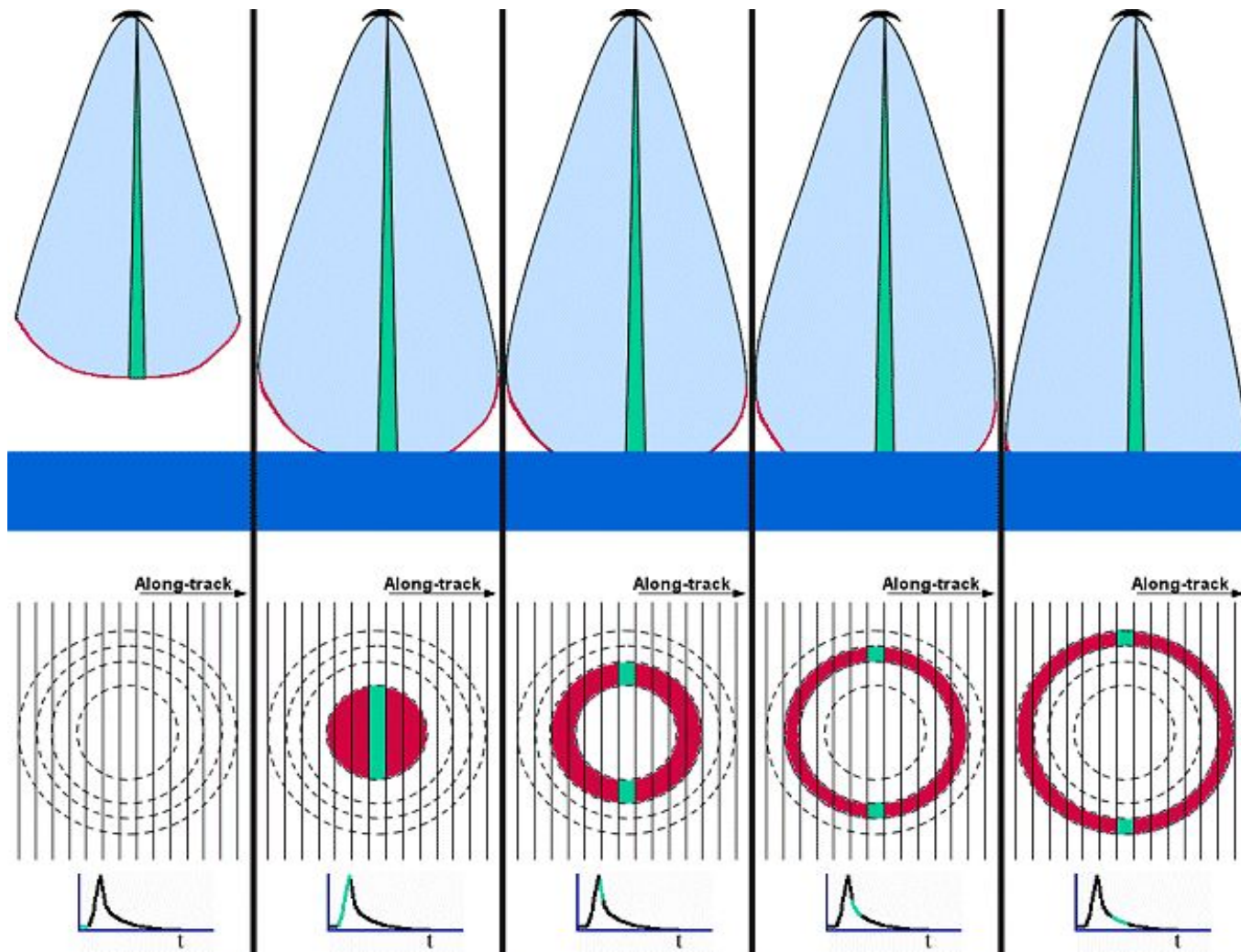


As a result of random distribution of the ocean wave facets at any instant, each individual return signal is very noisy, but averaging many successive pulses can reduce this.

12h of SAR Altimeters constellation (8 today)

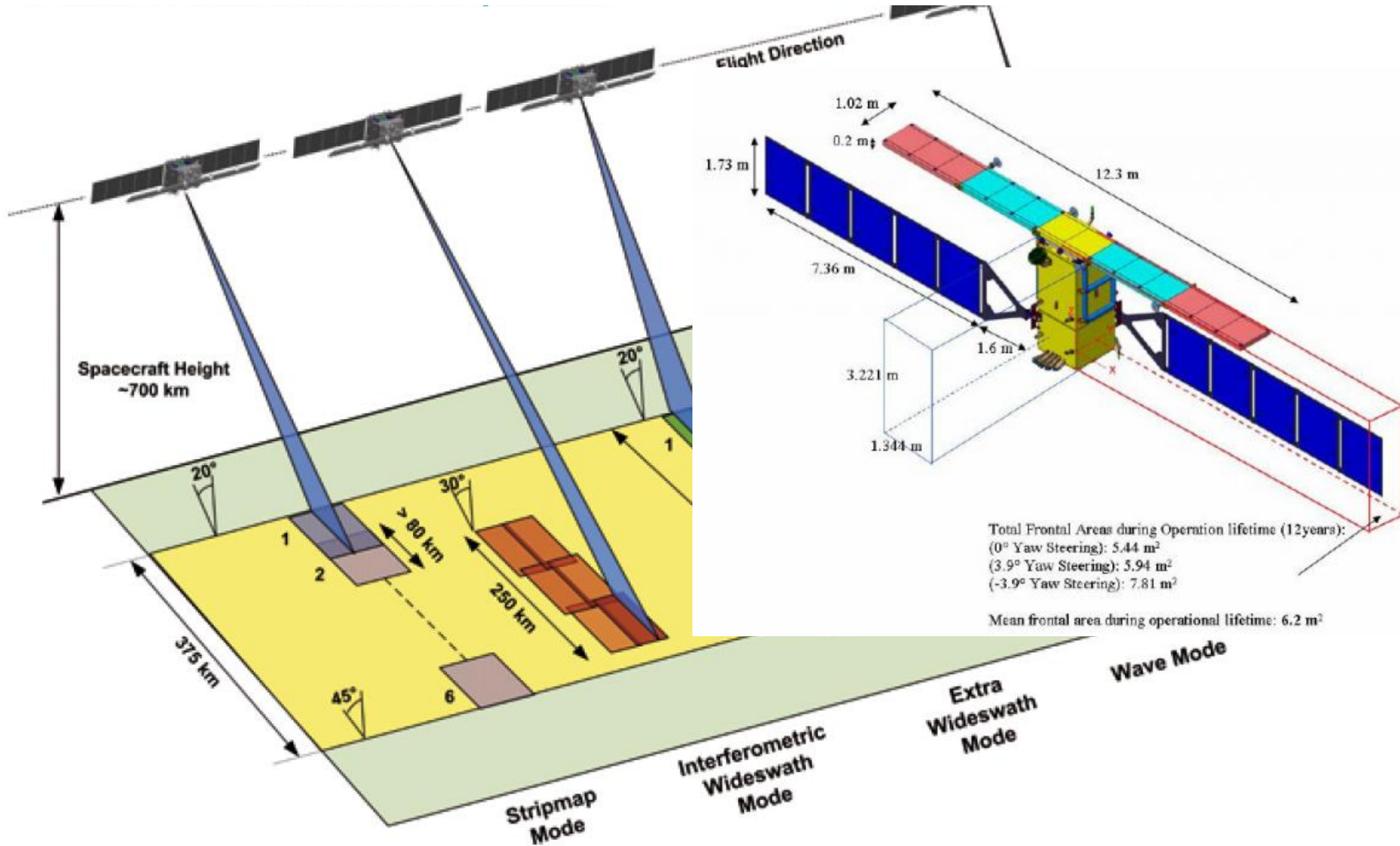


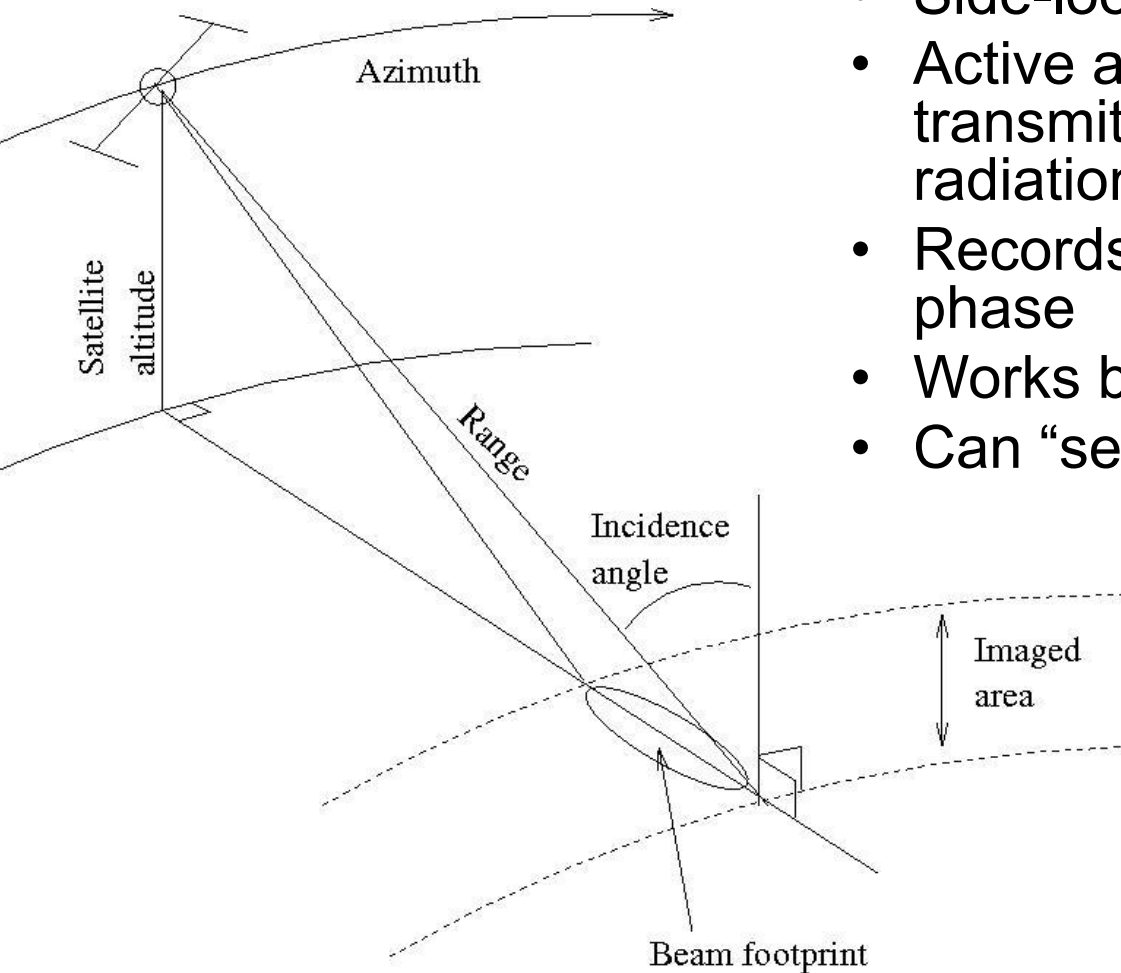
SAR Altimeters (New generation of altimeters)



Waves Remote Sensing by Synthetic Aperture Radar (SAR)

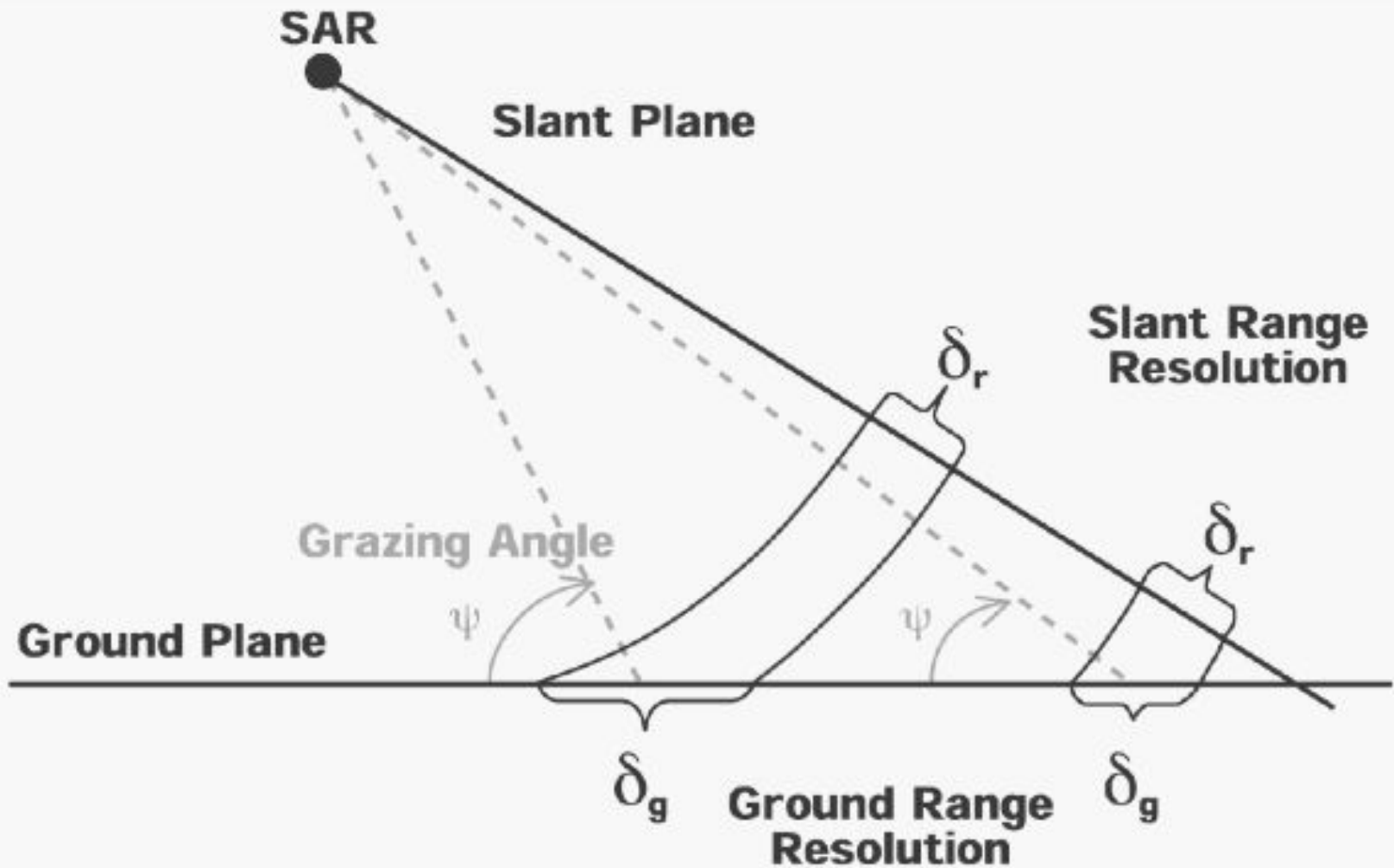
Sentinel-1 A&B (ESA,Copernicus)





- Side-looking
- Active antenna that transmits/receives electromagnetic radiation in VV, HH, VH, HV pol
- Records both signal amplitude and phase
- Works both day and night
- Can “see” through clouds

Ranges

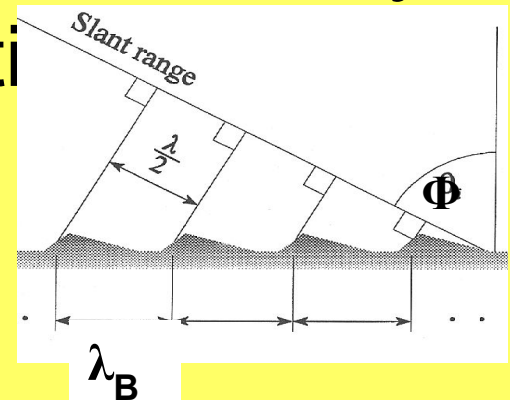


Scattering

At oblique incidence angles, the SAR backscattering arising from the sea surface is caused by surface waves of the order of the radar wavelength.

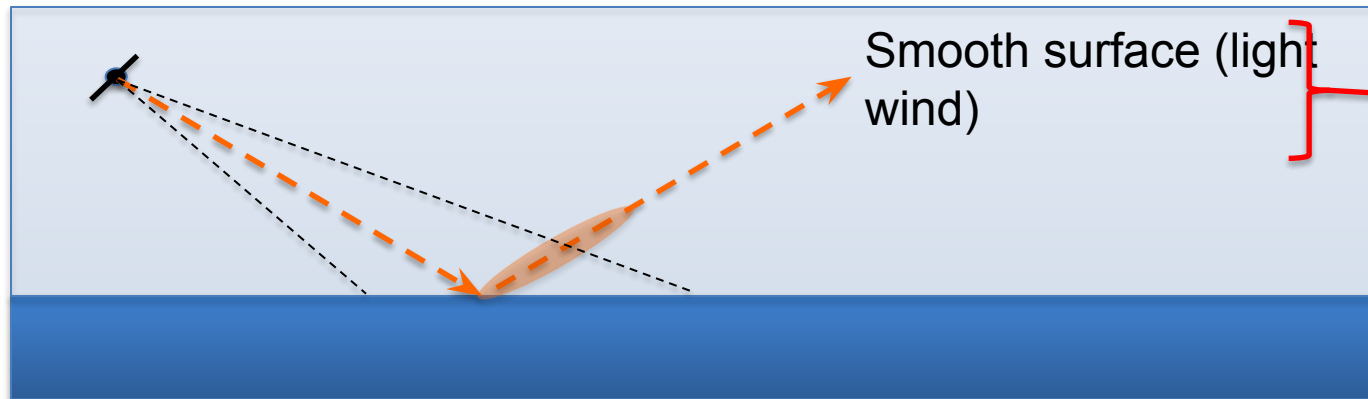
These waves are called “Bragg waves”. They obey the “Bragg resonance condition”

$$\lambda_B = \lambda_r / 2 \sin \Phi$$

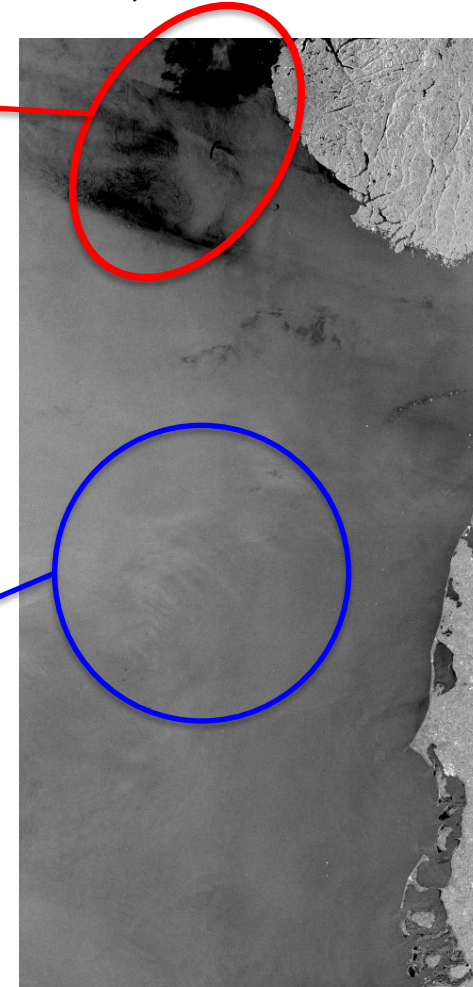
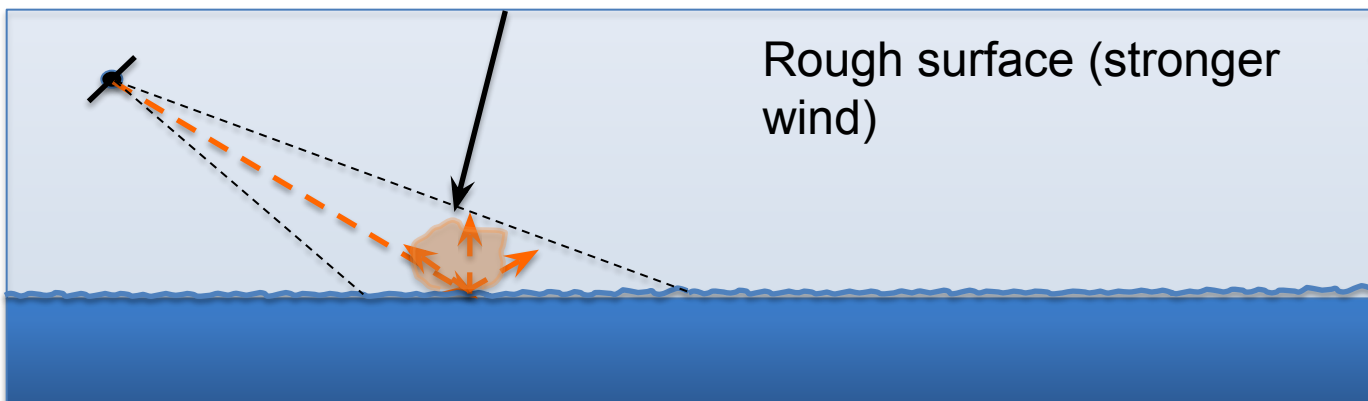


where λ_B = Bragg wavelength, λ_r = radar wavelength, and Φ = incidence angle

ASAR, 2.10.2011



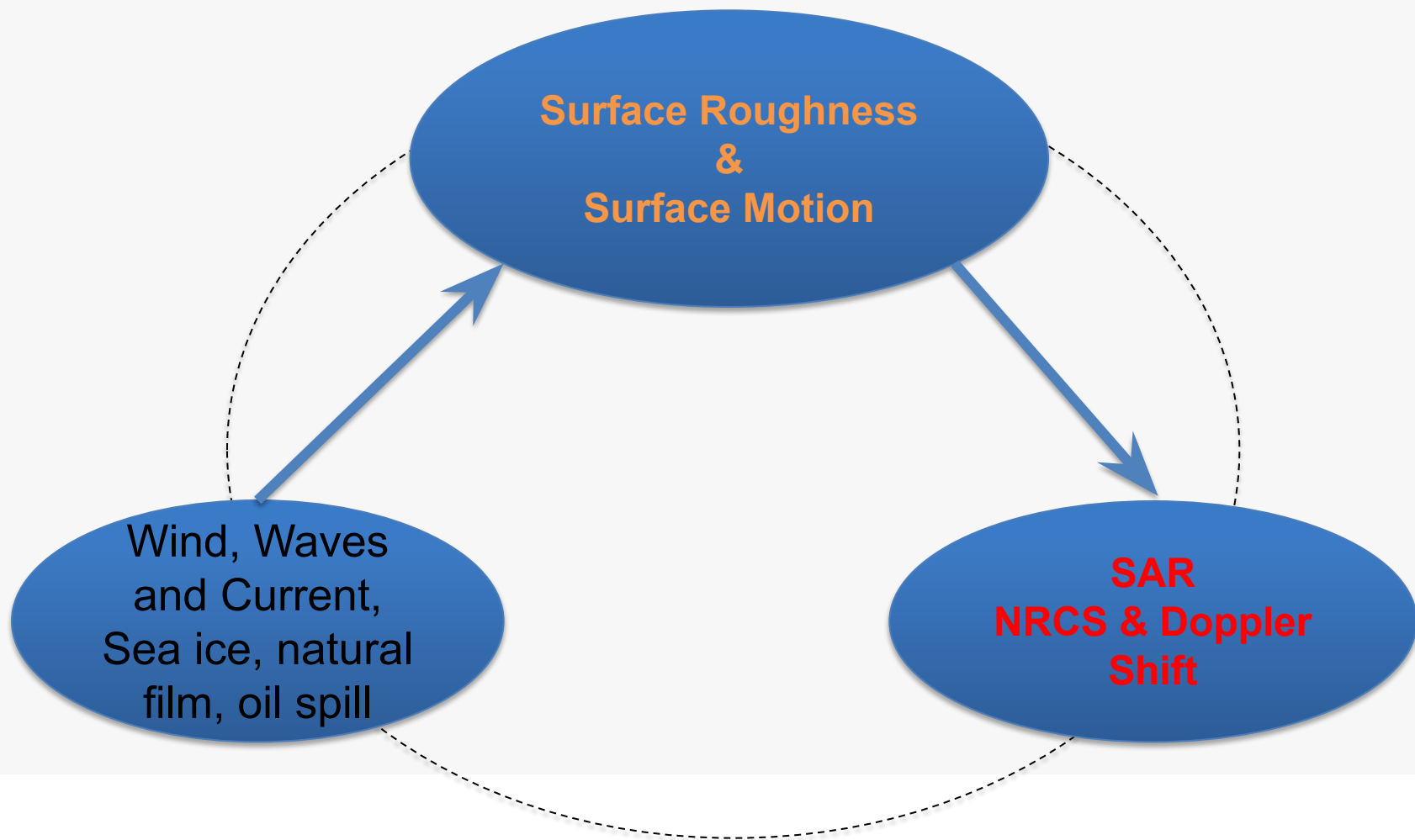
Enhanced
backscatter



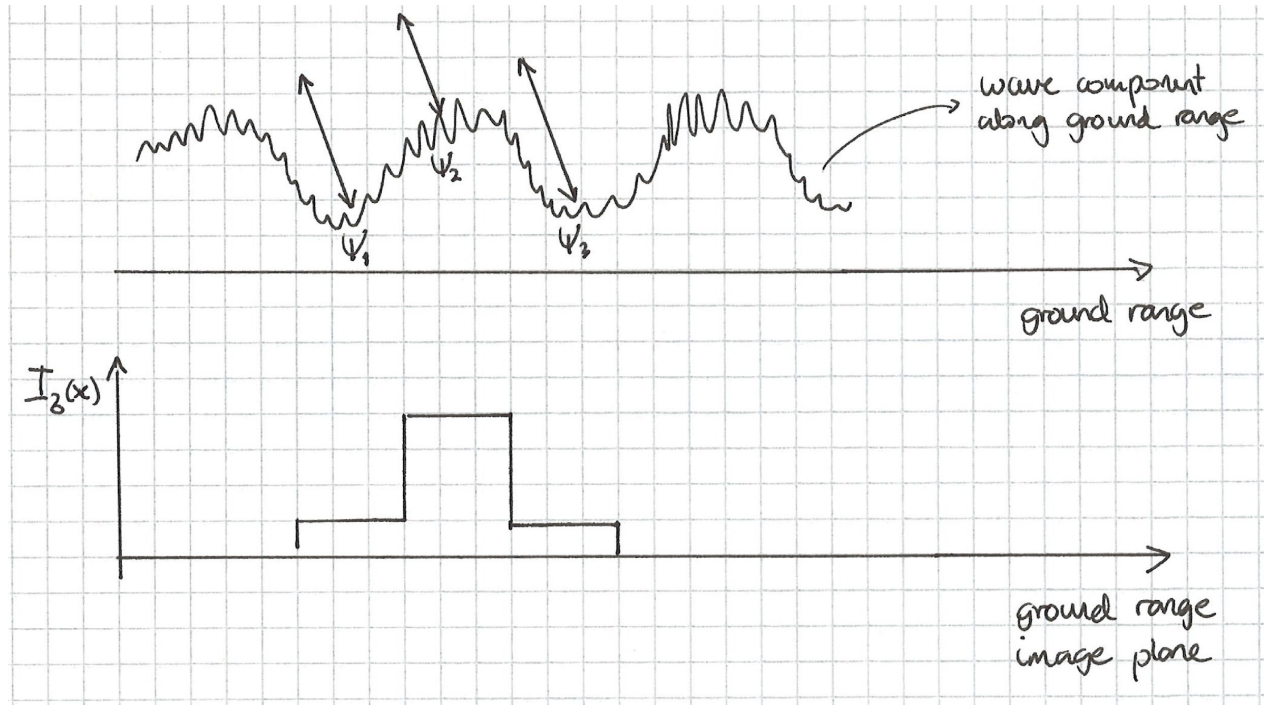
iceberg
The ocean surface roughness is influenced by **wind and waves, currents, surface slicks and sea ice** and is often different in open ocean versus coastal or ice covered regions due to fetch effects

The surface roughness is the source for the backscatter of the SAR signal.

The signal that arrives at the antenna is registered both in ***Amplitude and Phase***.



The Hydrodynamic Effect:



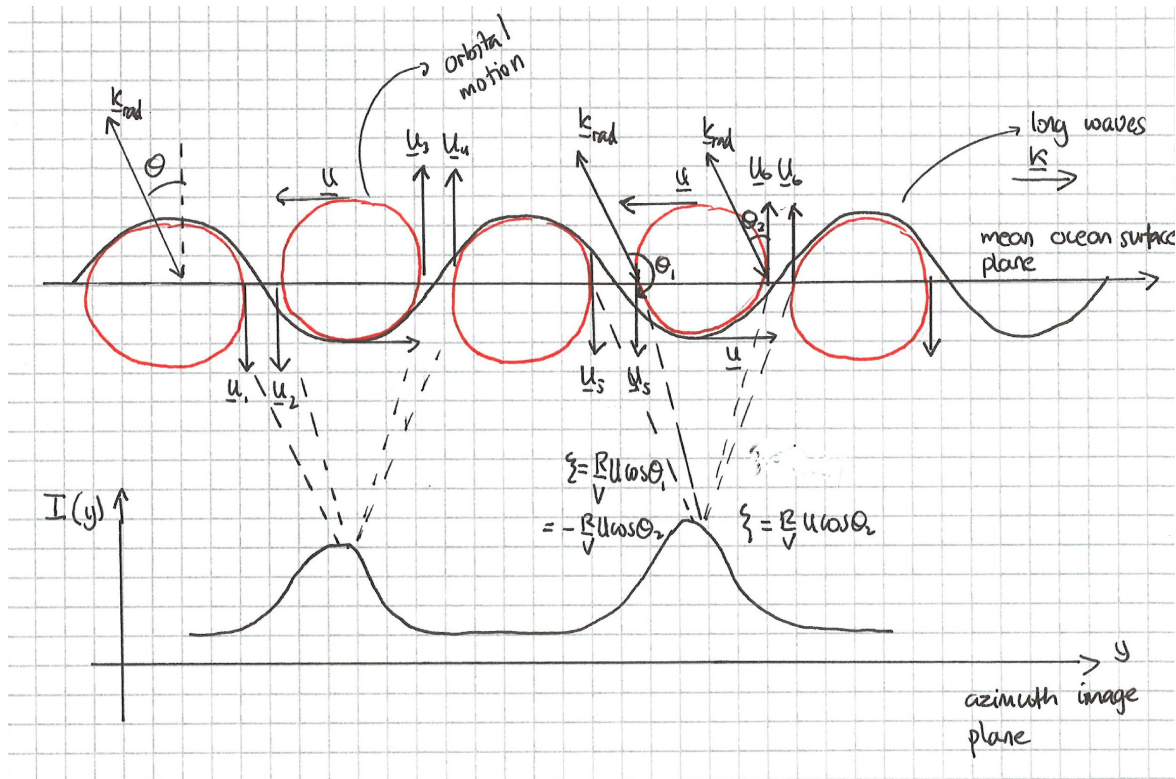
Backscattered energy density of radar resonant capillary waves spatially modulated by longer waves due to wave/wave and wave/wind interactions. The result is a spatial modulation in the backscatter image intensity

The Tilt Effect:

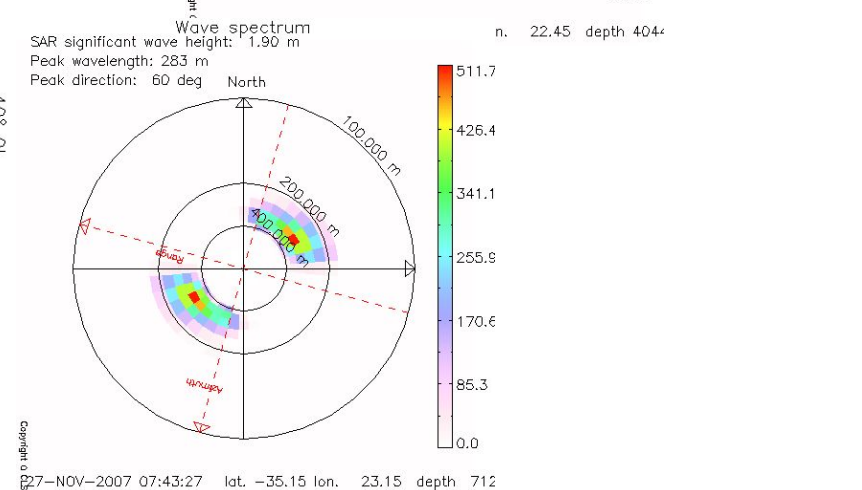
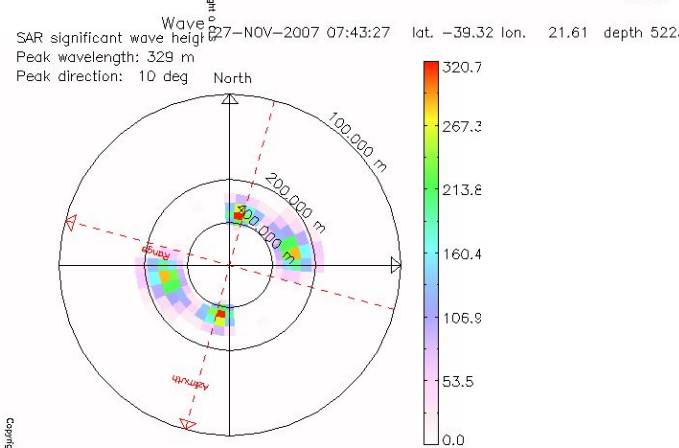
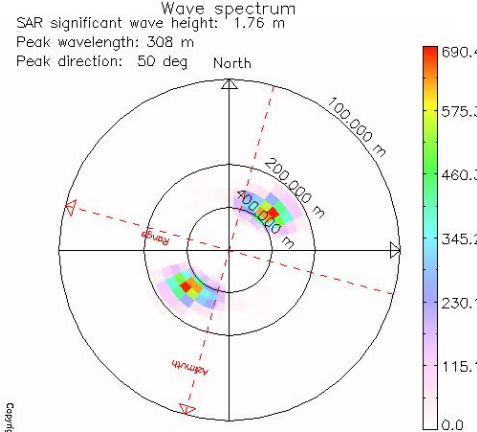
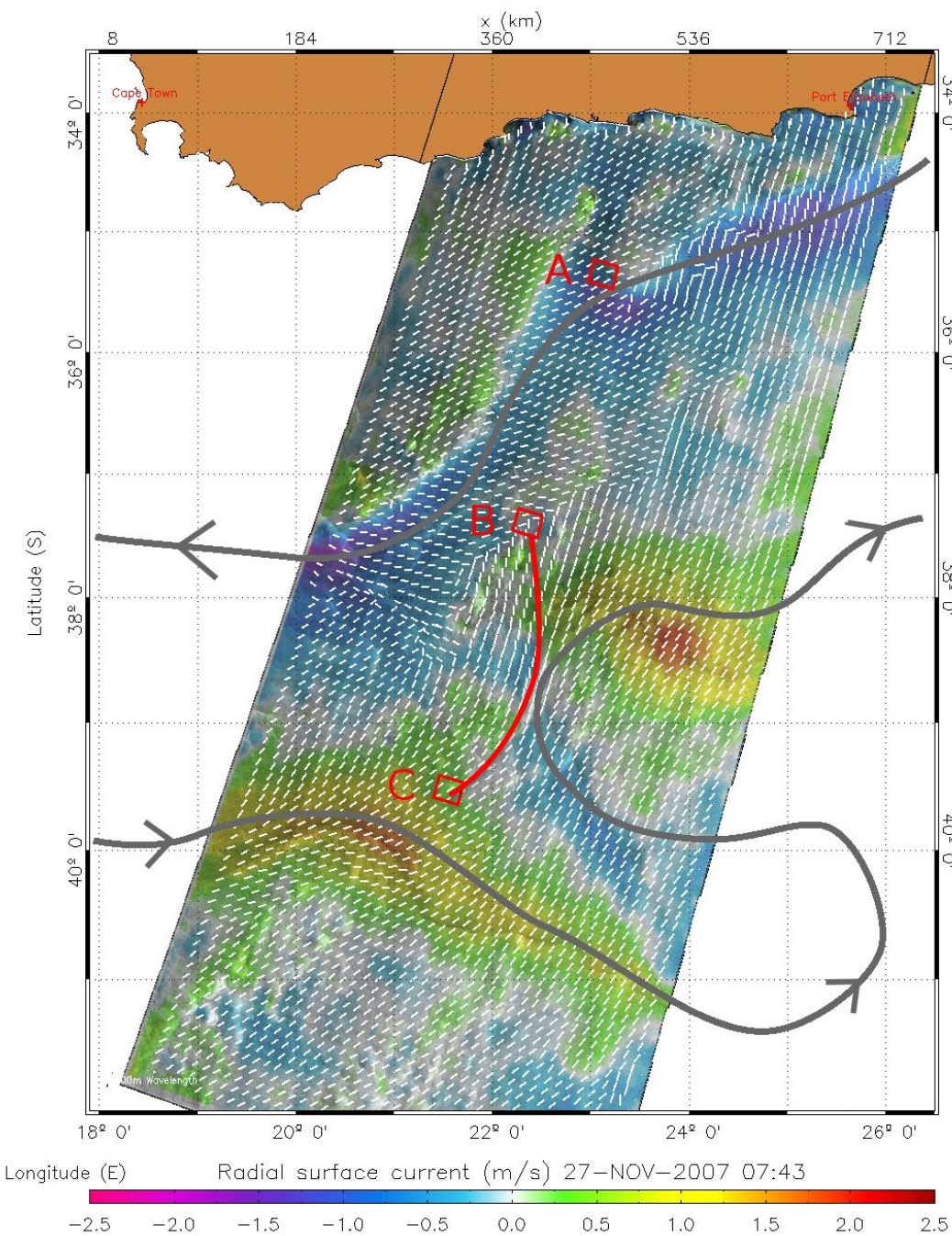


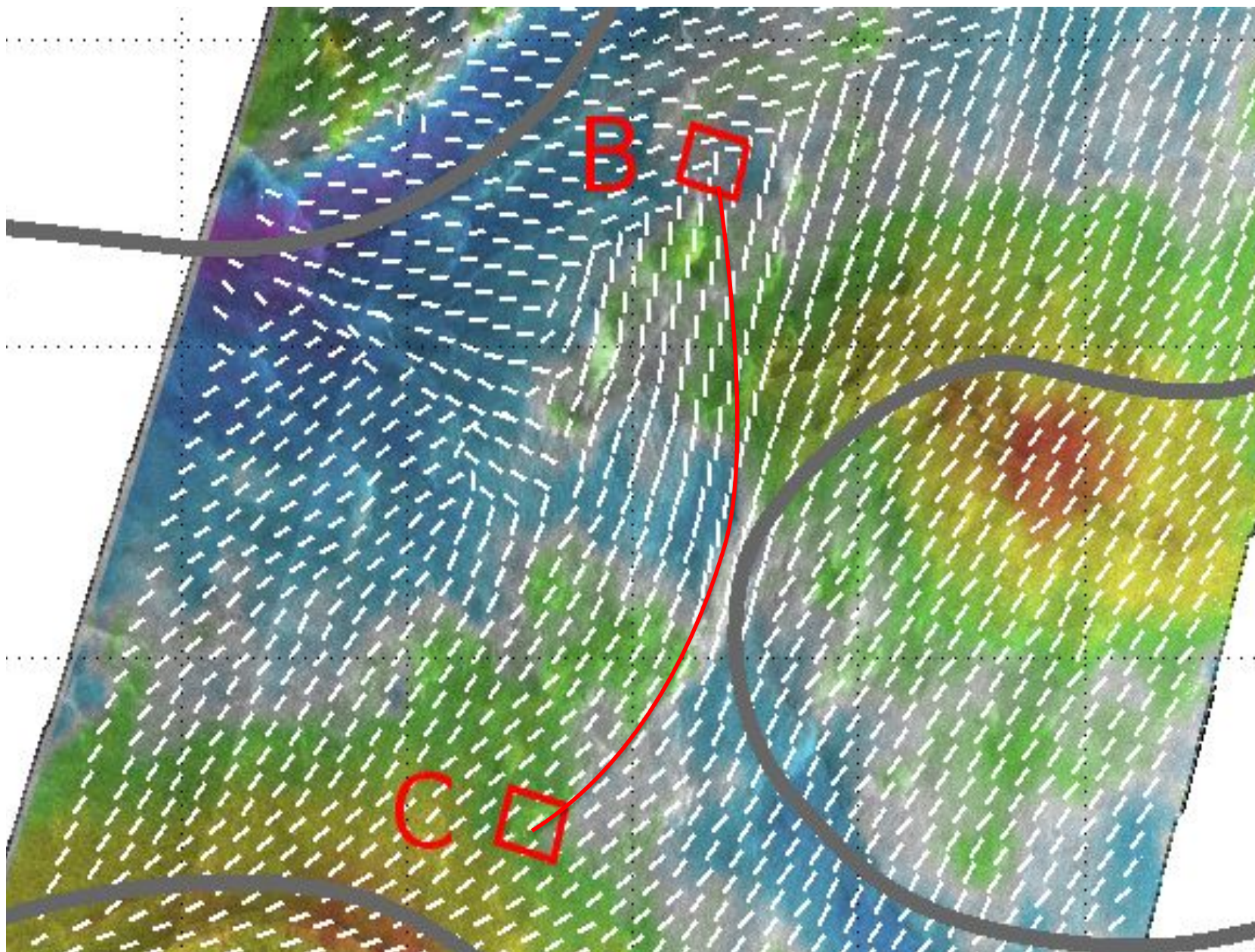
Backscattered energy density of radar resonant capillary waves spatially modulated by longer waves due to change in local incidence angle. The result is a spatial modulation in the backscatter image intensity

The Velocity Bunching (SAR) Effect:



The backscattered energy of the radar resonant capillary waves are shifted in azimuth by the Doppler effect due to the orbital motion of the scatterer on the longer waves. The result is a spatial modulation in azimuth in the SAR image intensity





Firework analysis : Principle

- Extraction of swell systems parameters from SAR WAVE MODE level2 products.
- Backward propagation to identify the swell origin (Storm source)
- Identification of all swell observations relative to a given Storm source.
- Determination of the propagation path by forward and backward propagation between observations (dispersion relationship).

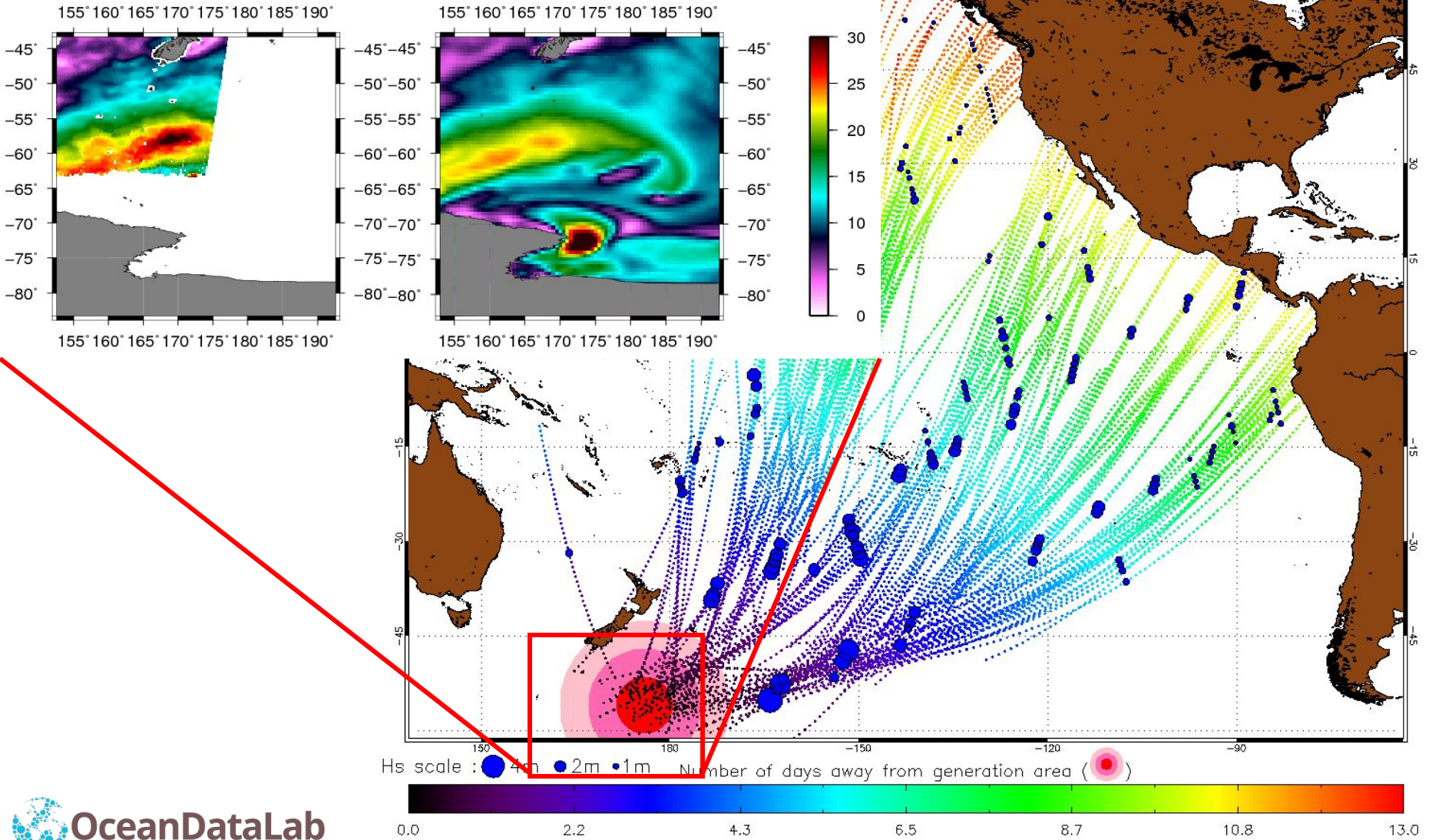
Propagation of 15-16s swell from July 8 to July 20, 2004

Wind speed: 32.9 m/s

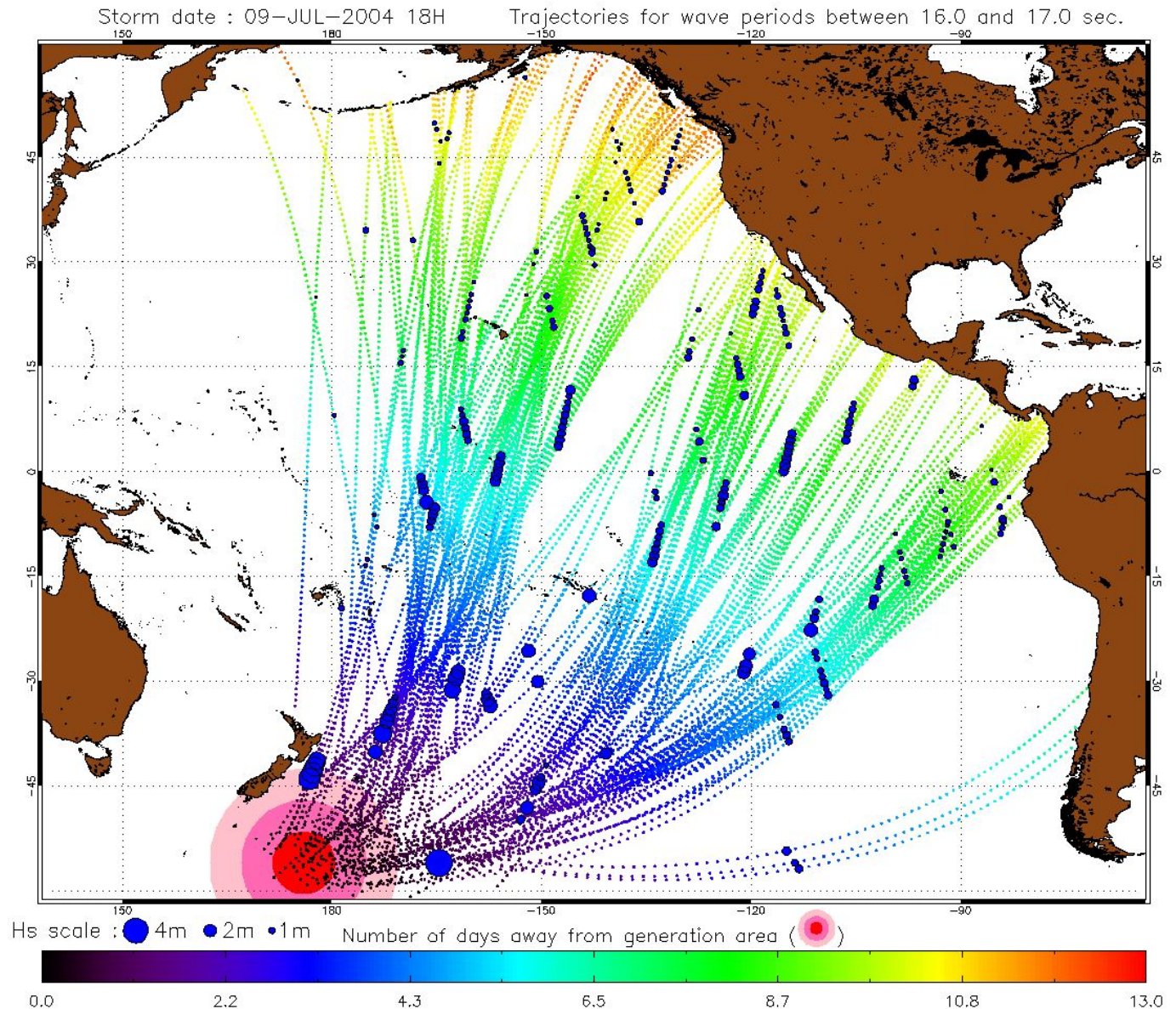
Swath date: 09/07/2004 06:26

Model date: 09/07/2004 06:00

stories for wave periods between 15.6 and 16.0 sec.



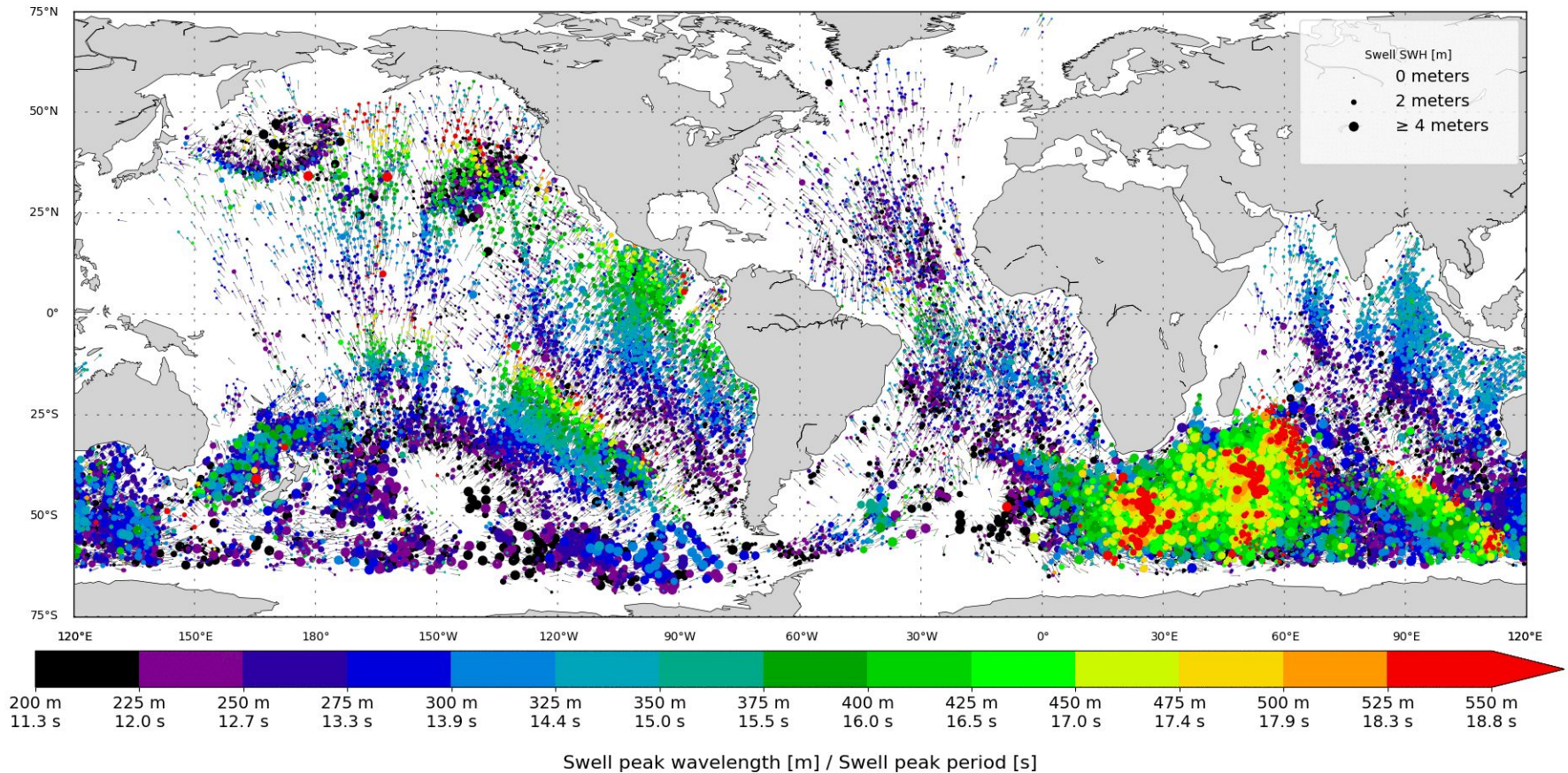
Propagation of 16-17s swell from July 8 to July 20, 2004



Today with Sentinel-1A&B Wave mode

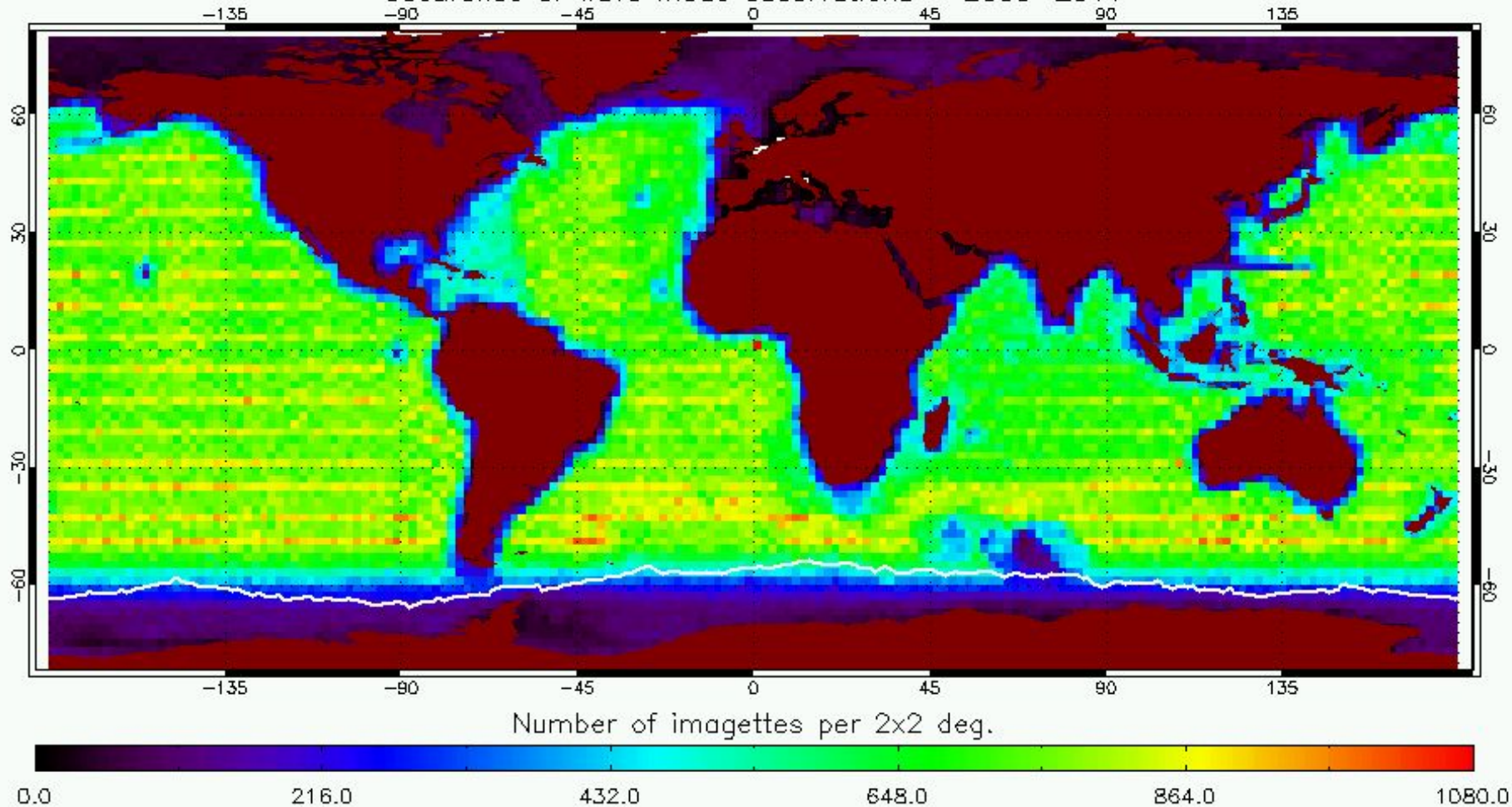
E.U. Copernicus Marine Service Information

Fireworks using S1A/S1B - 03 Jun 2022 00H UTC

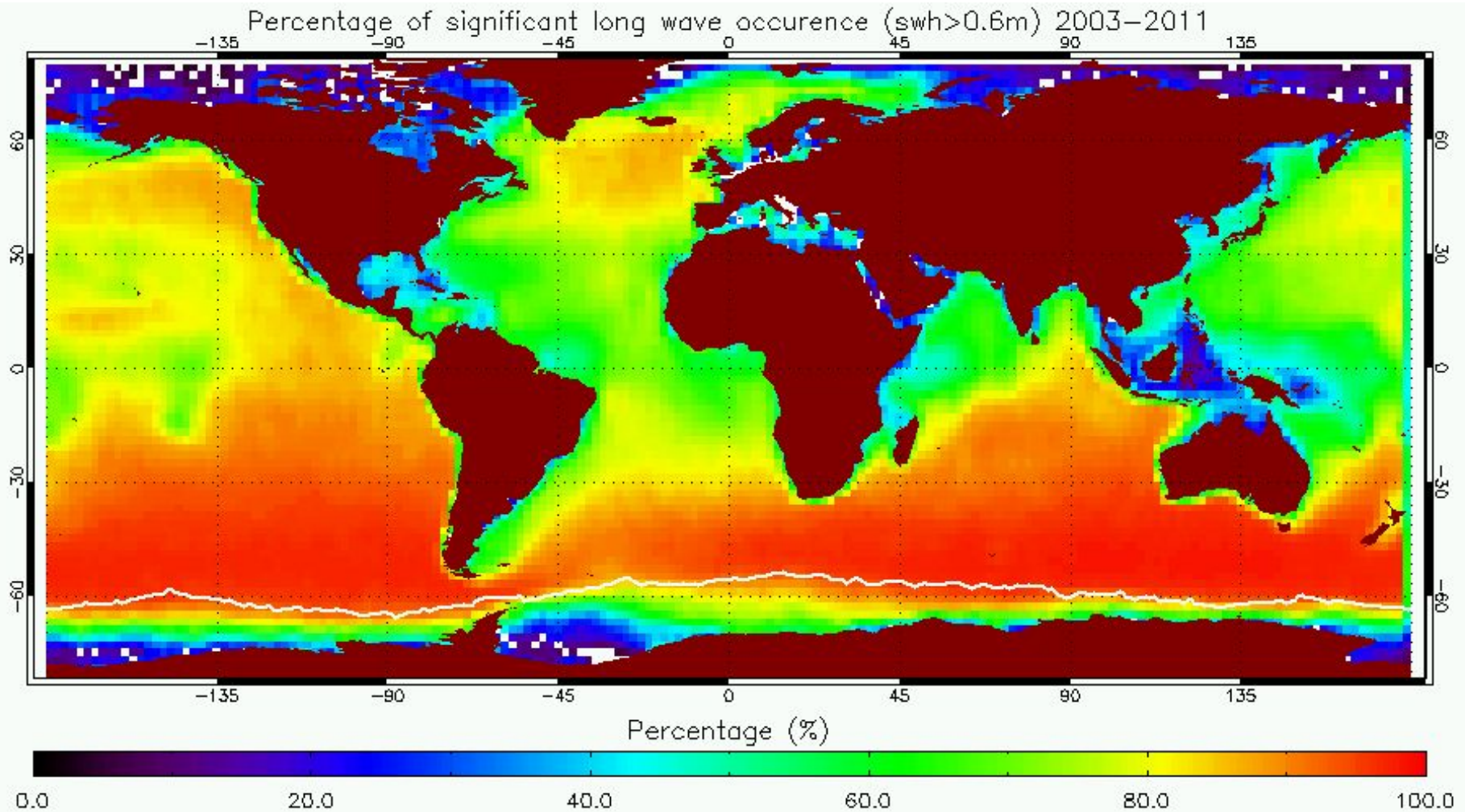


Statistical analysis ENVISAT ASAR Wave mode

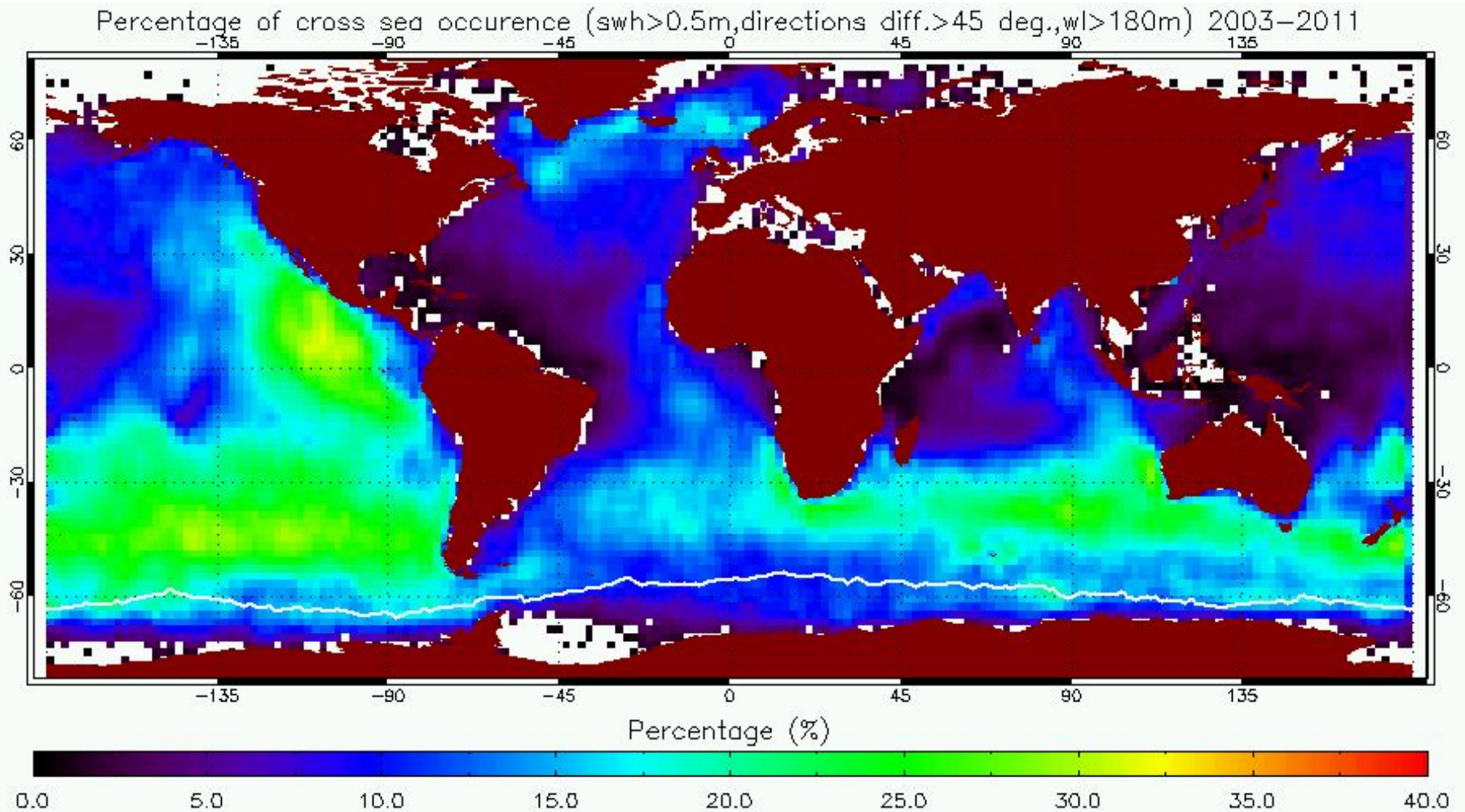
Occurrence of wave mode observations = 2003–2011



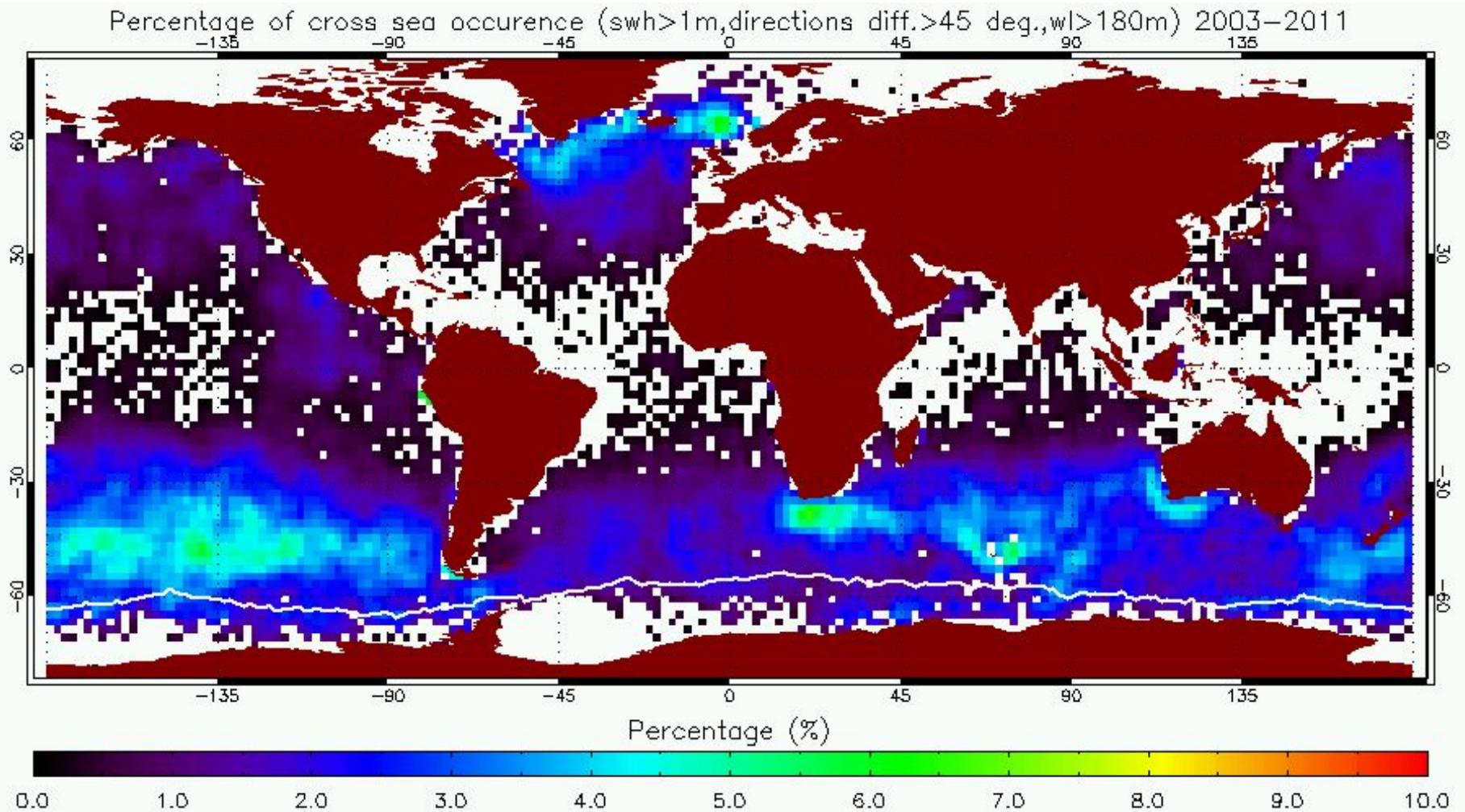
Percentage of detected long waves/swell



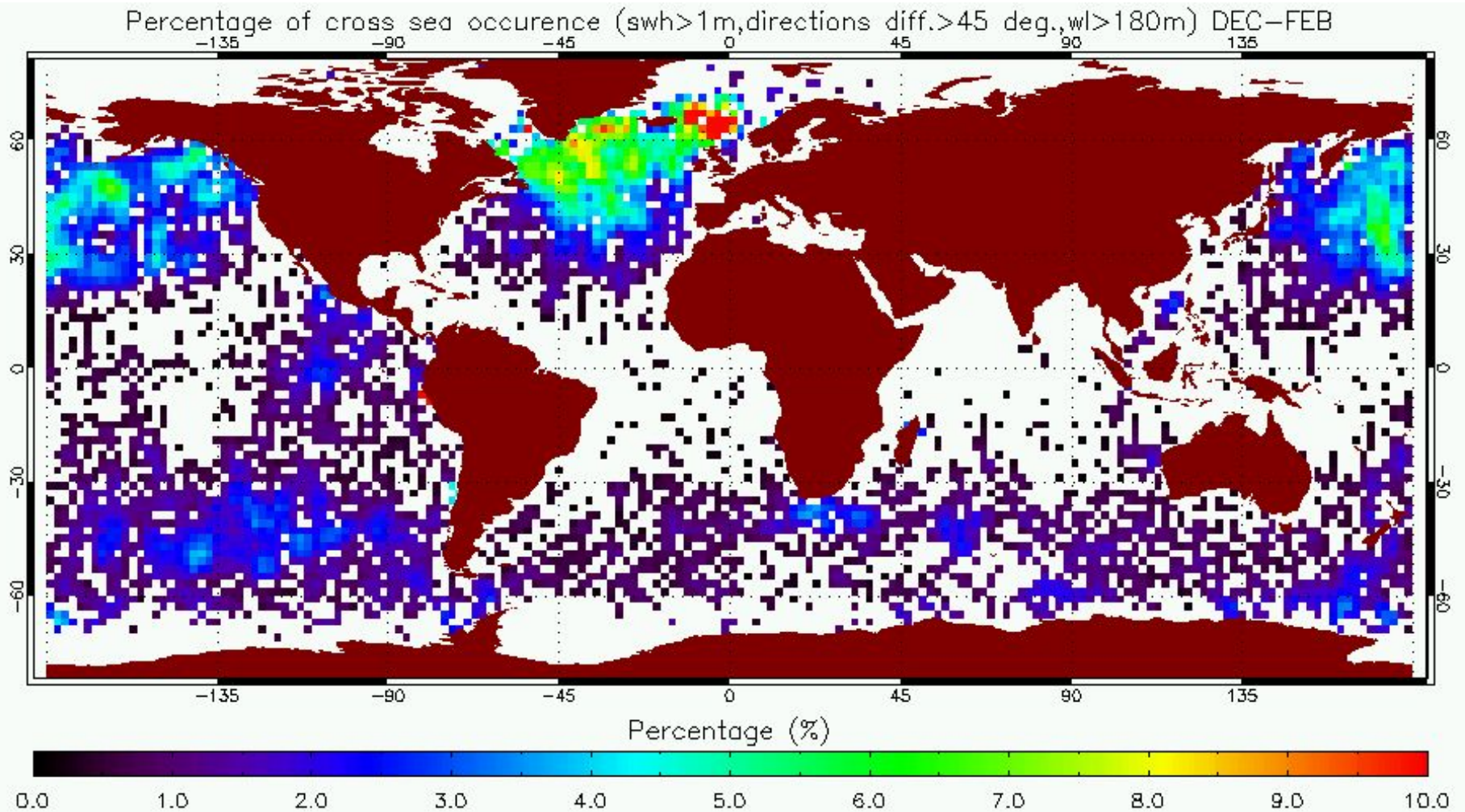
Percentage of detected cross swells



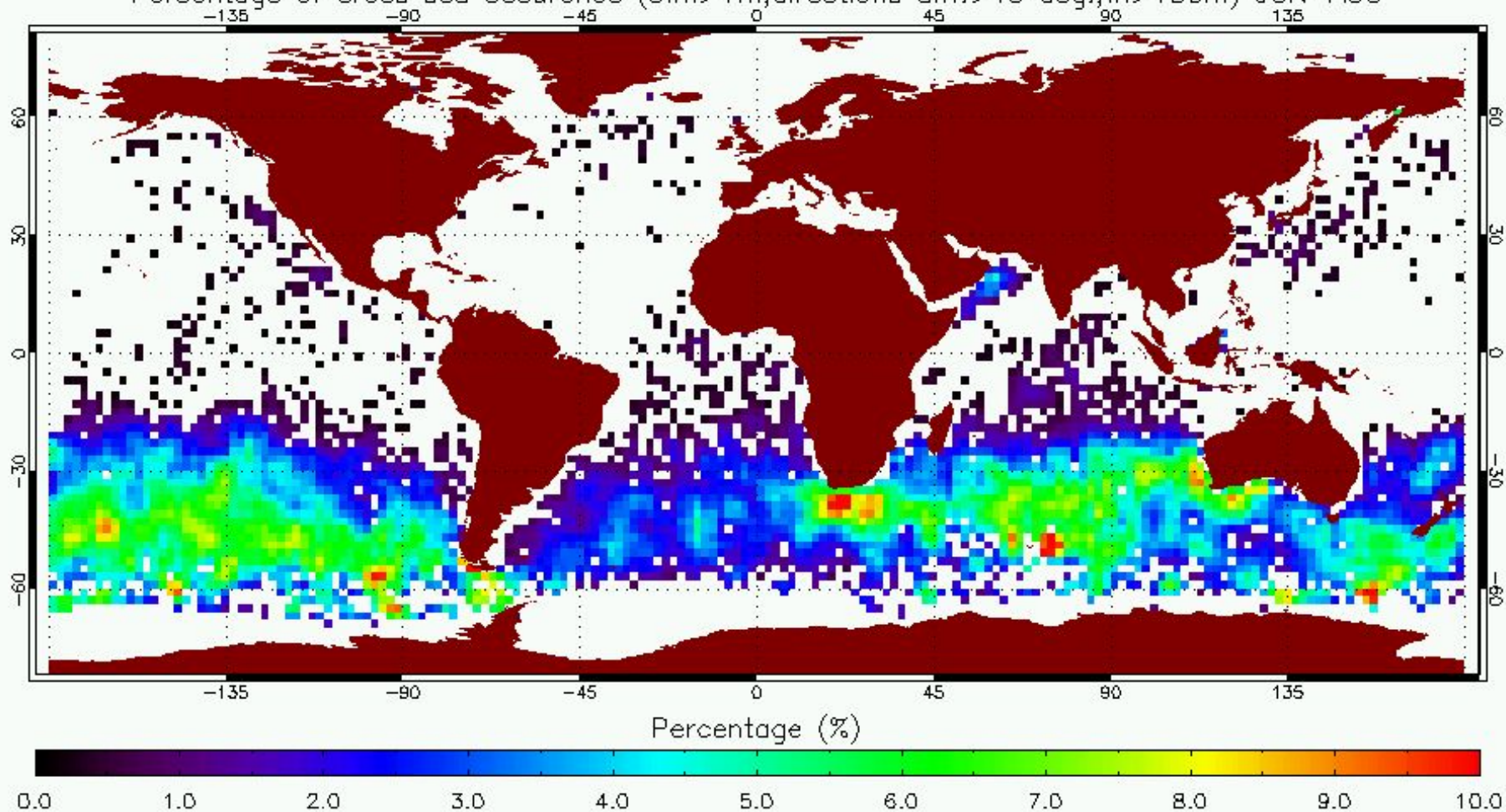
Percentage of detected cross swells > 1m Hs



Seasonal variability of cross swells > 1m Hs

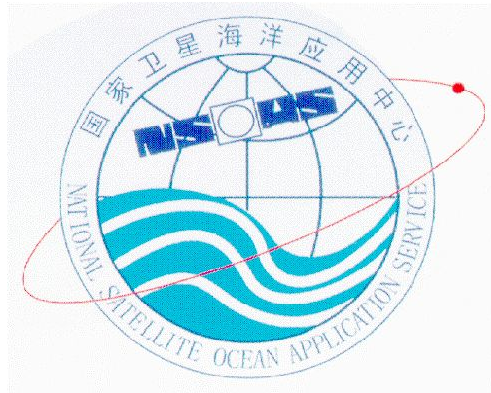


Percentage of cross sea occurrence (swh>1m,directions diff.>45 deg.,wl>180m) JUN–AUG



CFOSAT (launched in 2018)

Wind and wave observations from space: A French-Chinese mission



Scientific PIs:

Lotfi Aouf (France): Météo-France (previously Danièle Hauser)

Liu Jianqiang (China): NSOA

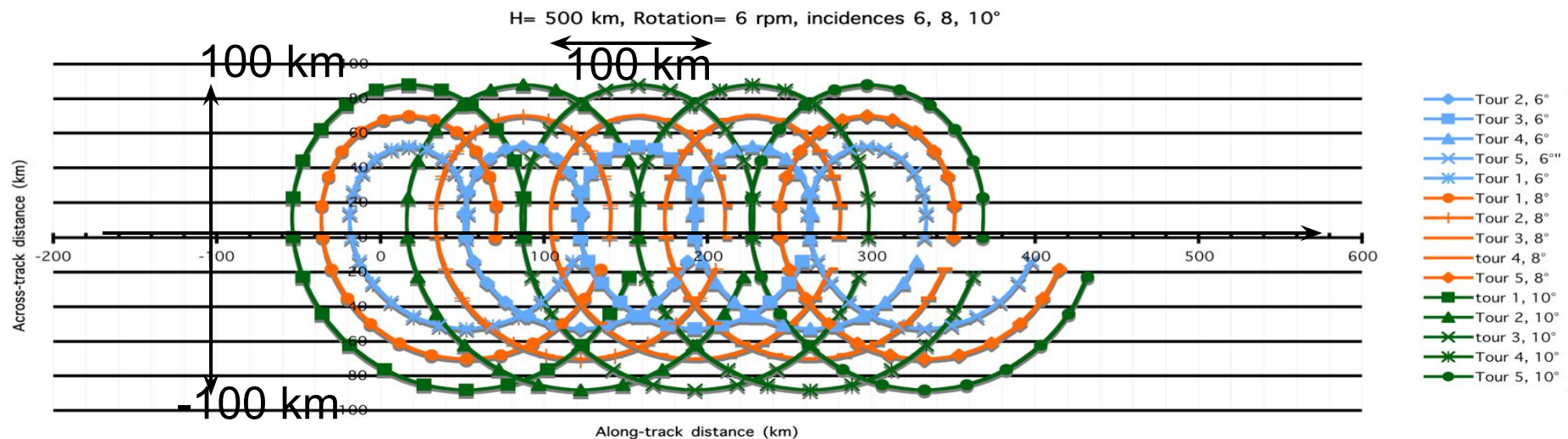
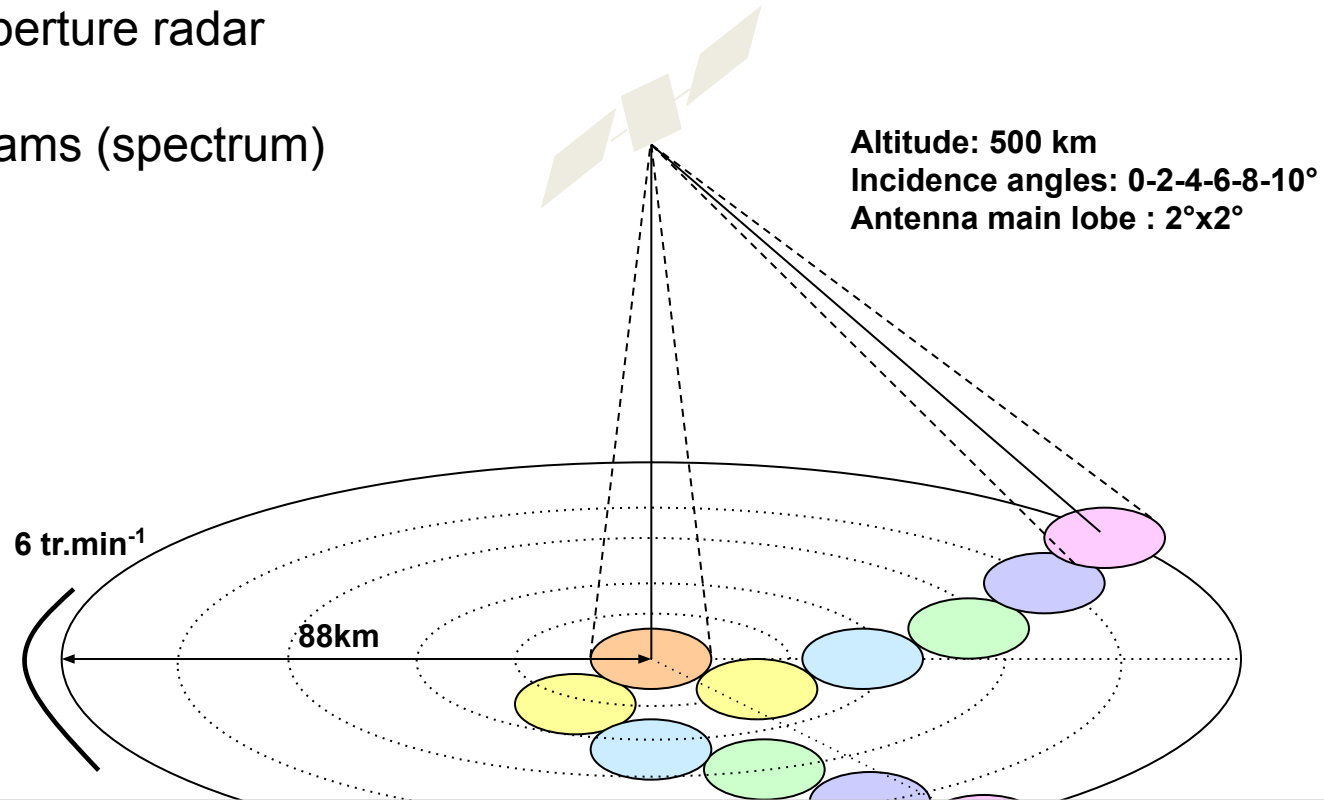
and the Mission and Technical Team

SWIM instrument

SWIM: real aperture radar

1 nadir beam

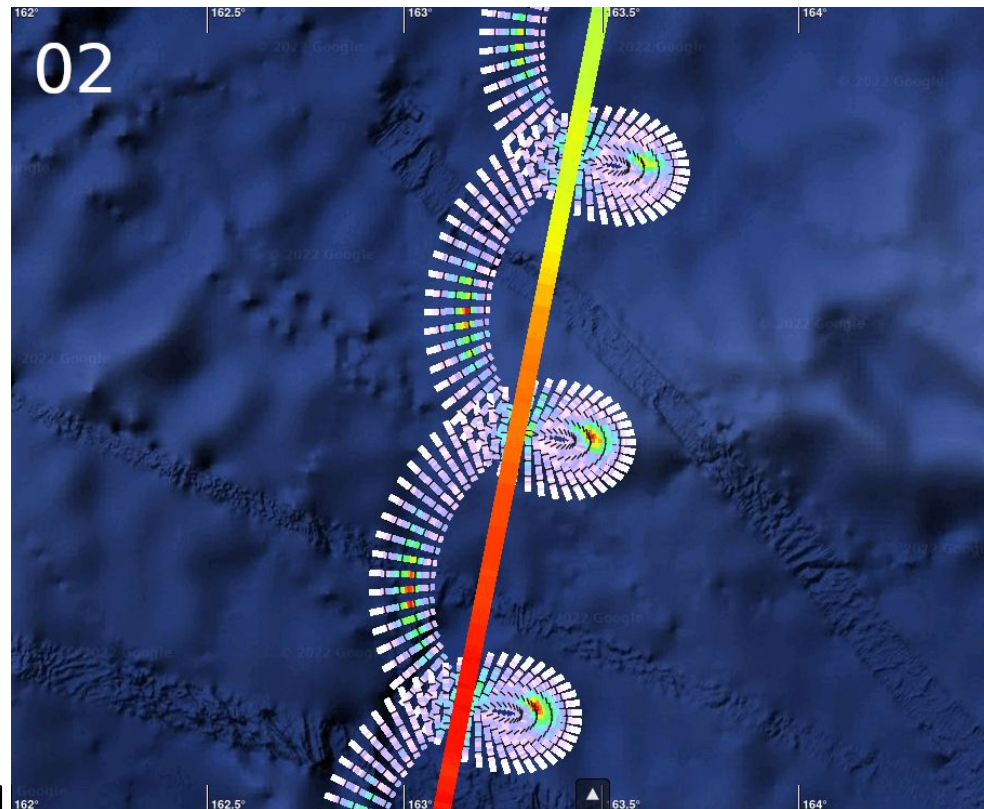
5 off-nadir beams (spectrum)



SWIM instrument

- SWIM (Surface Wave Investigation and Monitoring) measures the ocean surface wave related modulations in Ku band using a rotating instrument
- In nominal macrocycle mode, it provides a directional 1D wave spectra
 - every ~ 7 degrees
 - for each of its 5 beams at 2, 4, 6, 8, and 10° incidence angle
 - resulting in a very special cycloid ground footprint geometry.

Example of 1D raw spectra
projected over range footprint
With nadir Hs

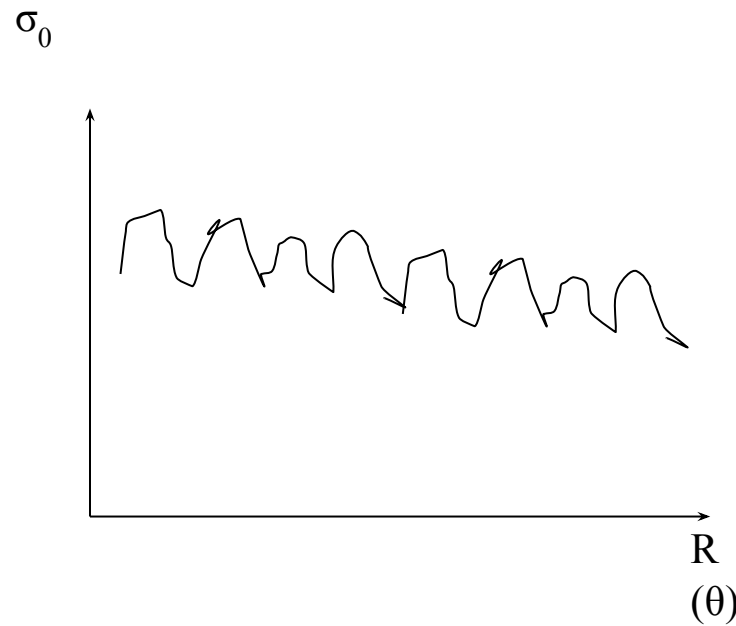
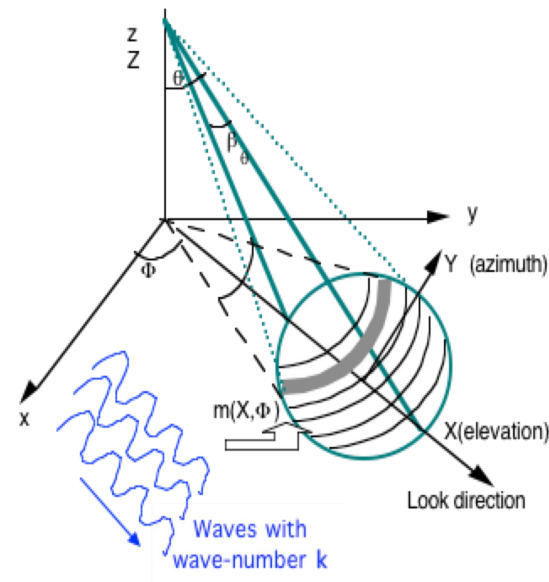


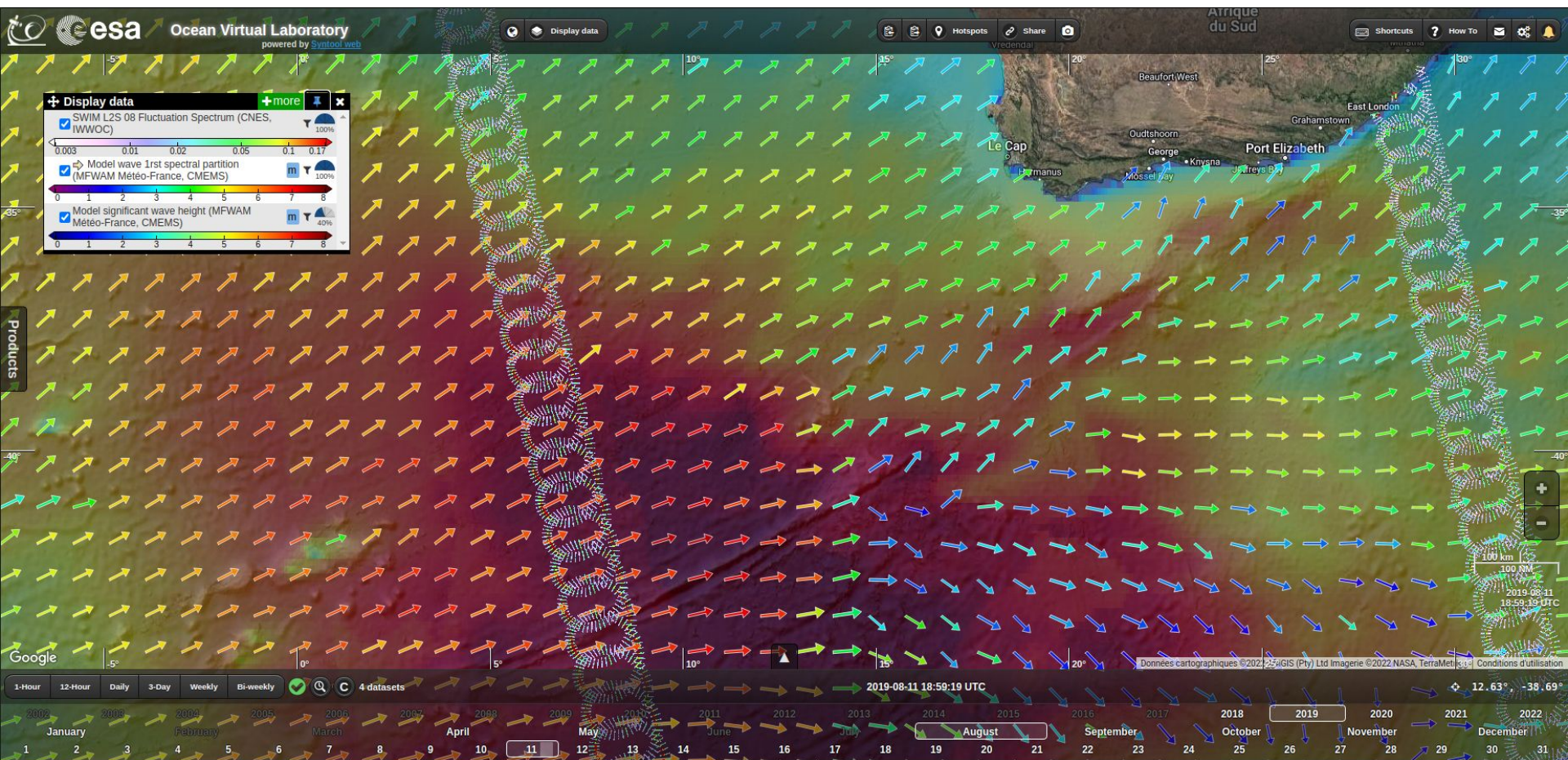
-Near Real time transmission of data (within 3 hours)

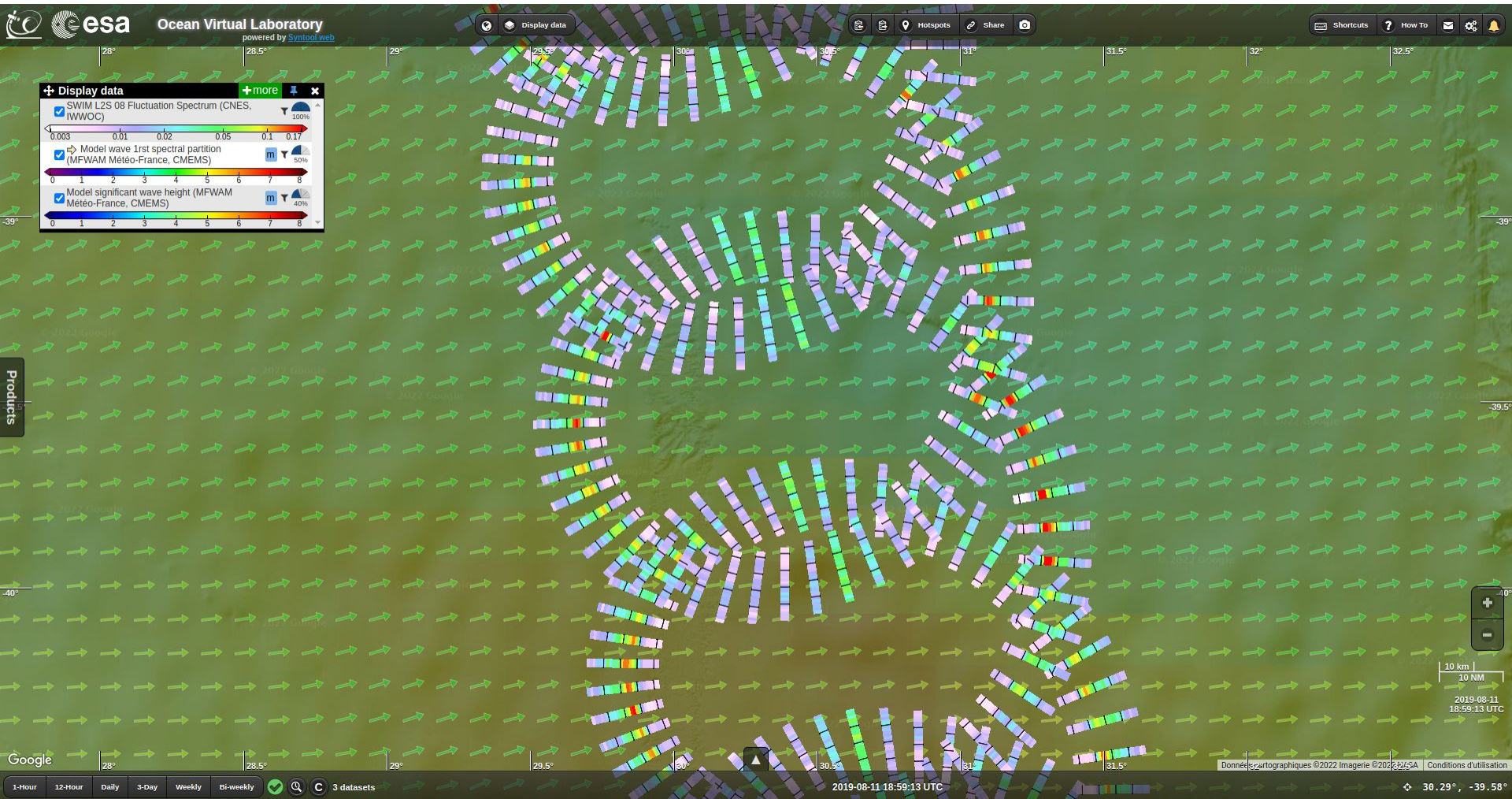
Principle for estimating wave spectra

- In each azimuth direction, the **normalized radar cross-section** σ_0 is modulated by the **tilt of the long waves**

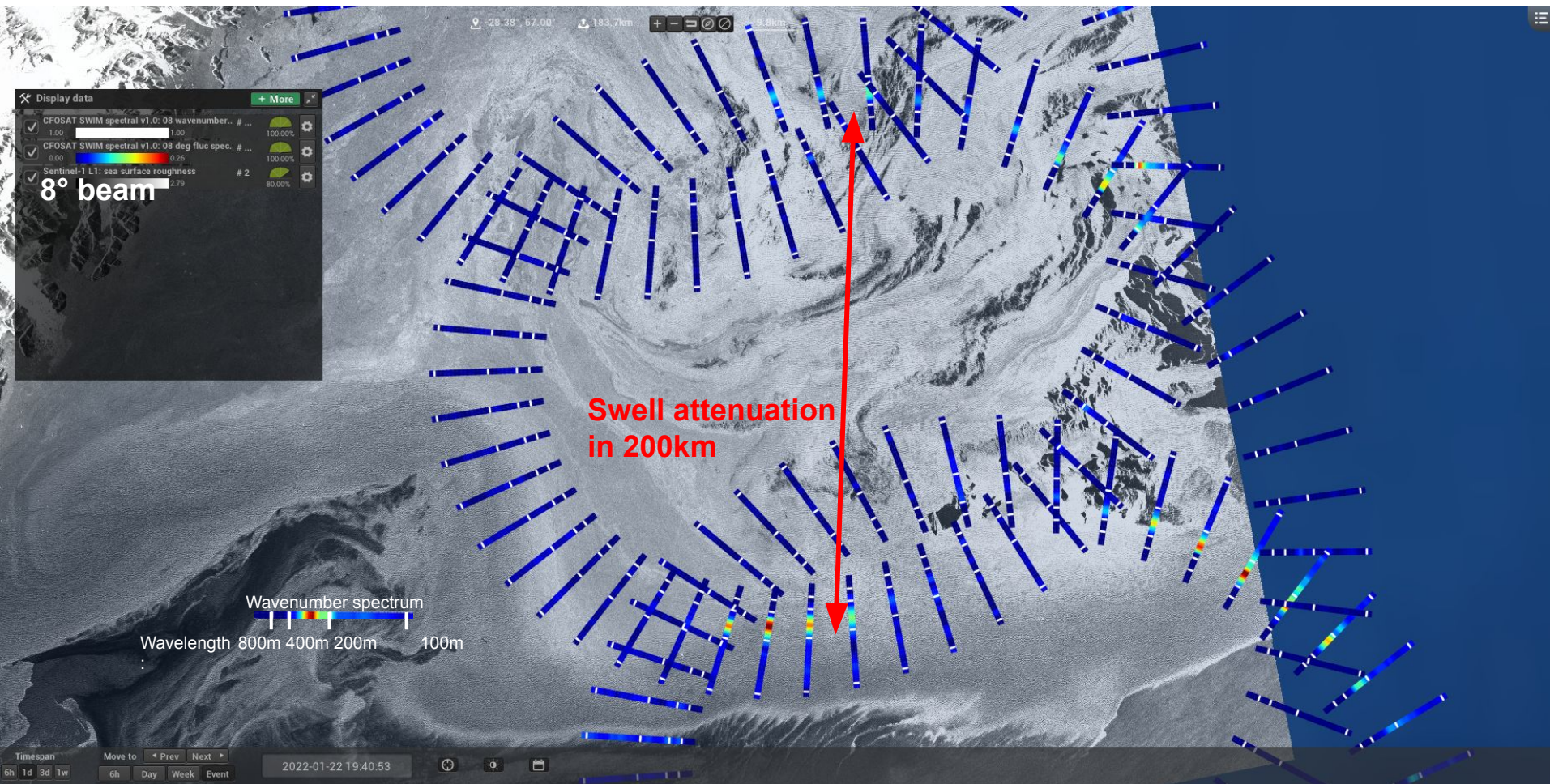
=> Measurement of **these modulations** $m(x, \phi)$, calculation of their **spectrum** $P_m(k, \phi)$



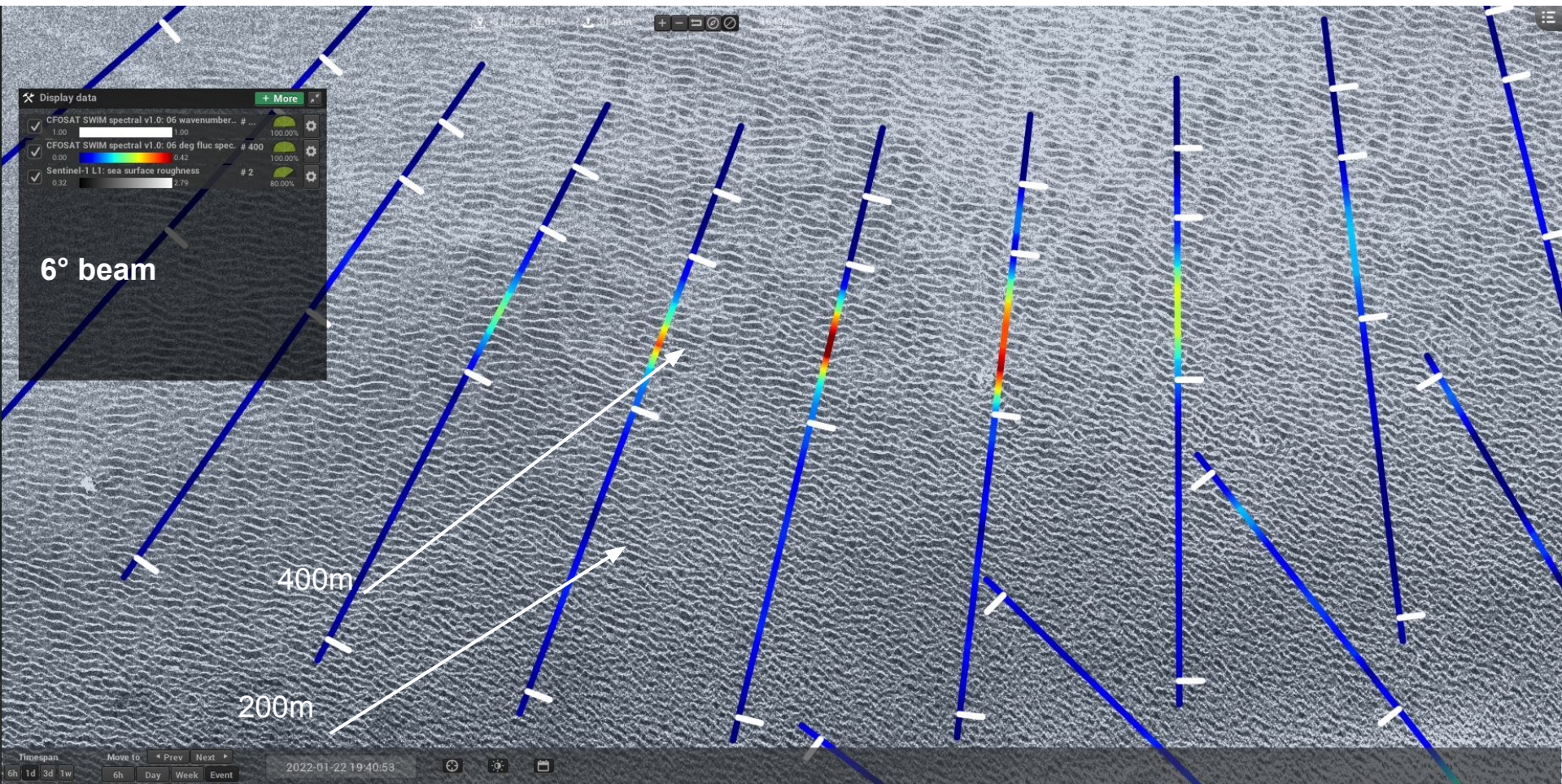




5. Last sensor: CFOSAT's SWIM instrument



5. Last sensor: CFOSAT's SWIM instrument



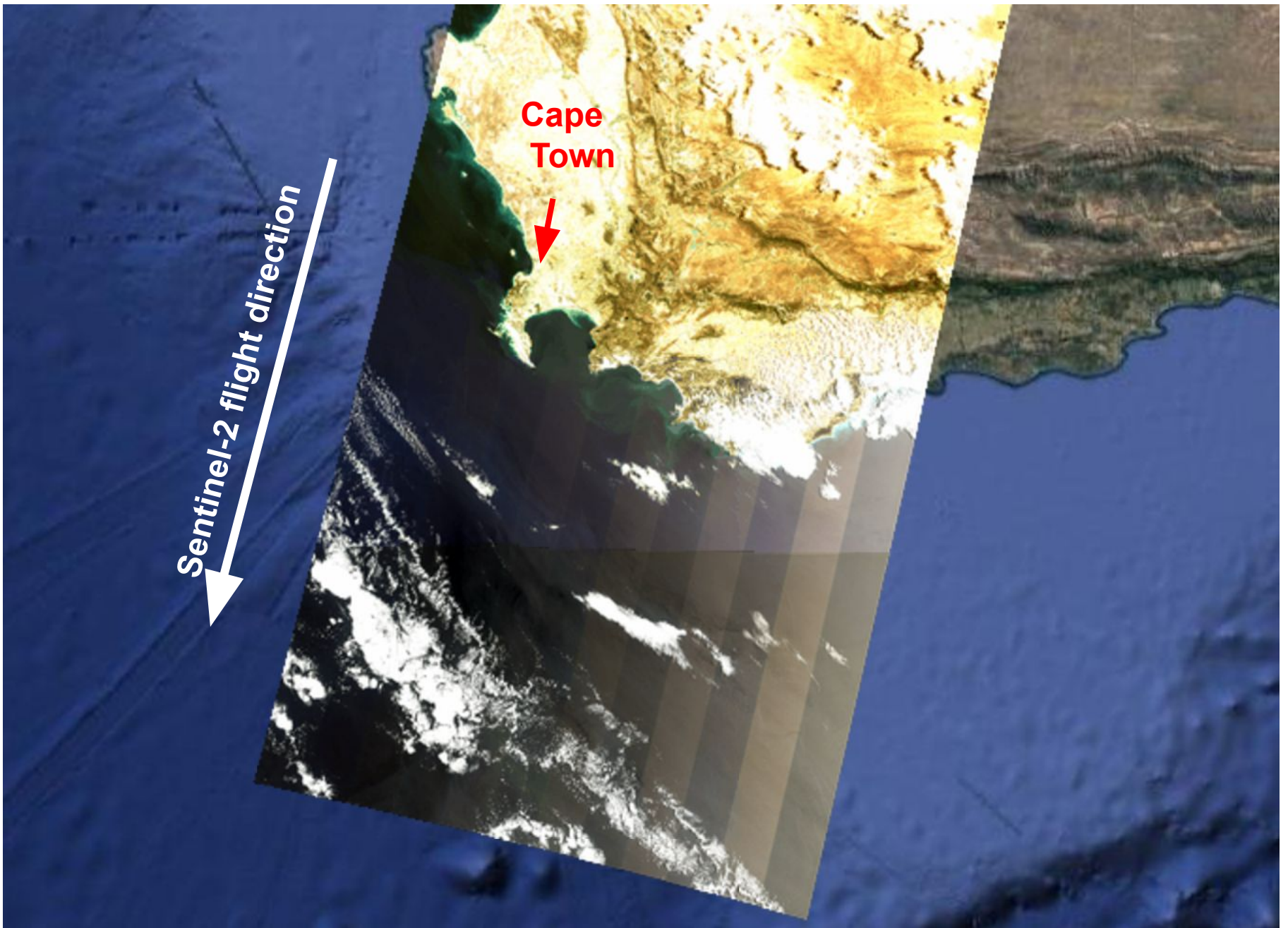
S1A/B 20210127T00

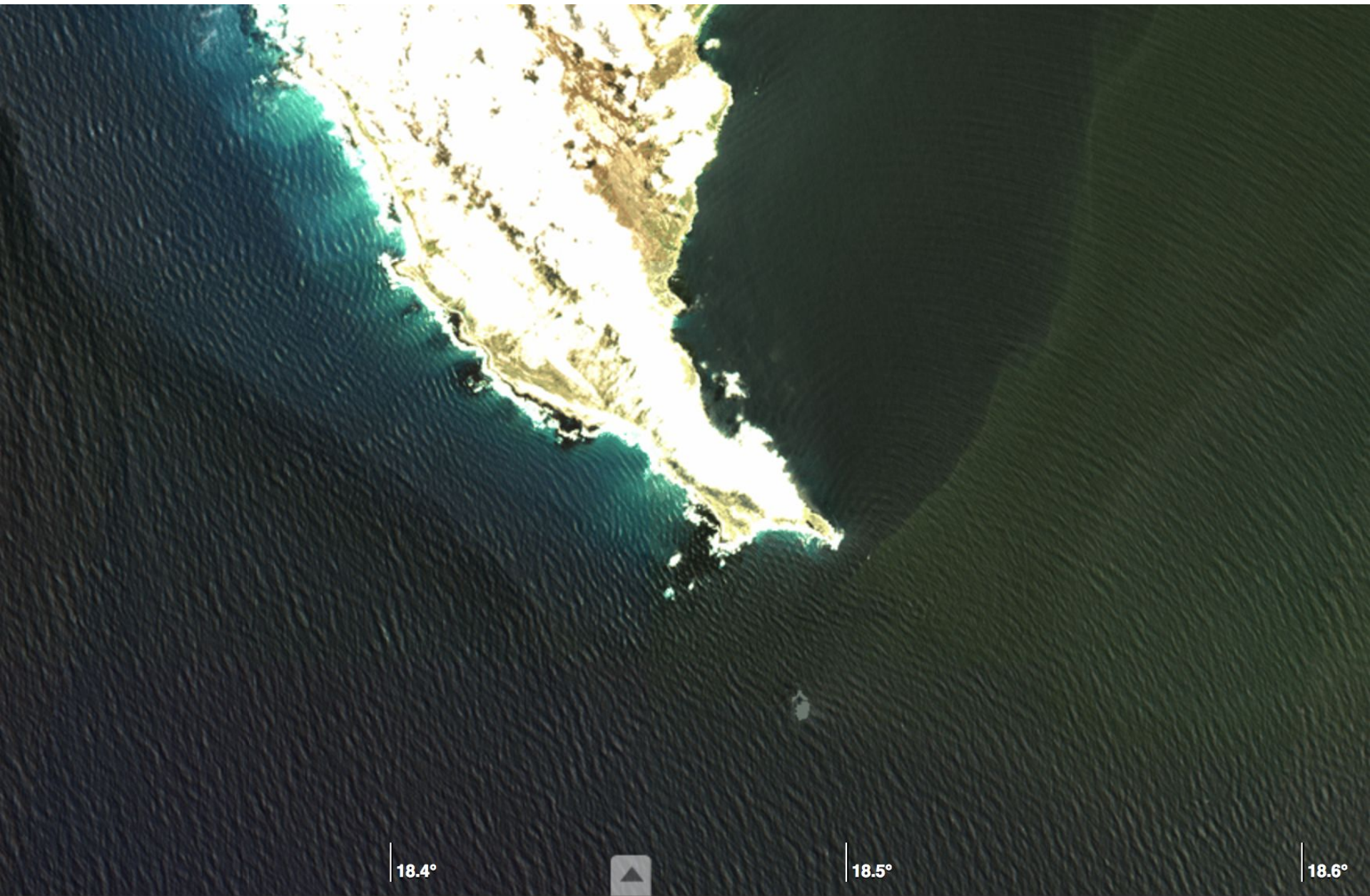


SWIM 20210127T00

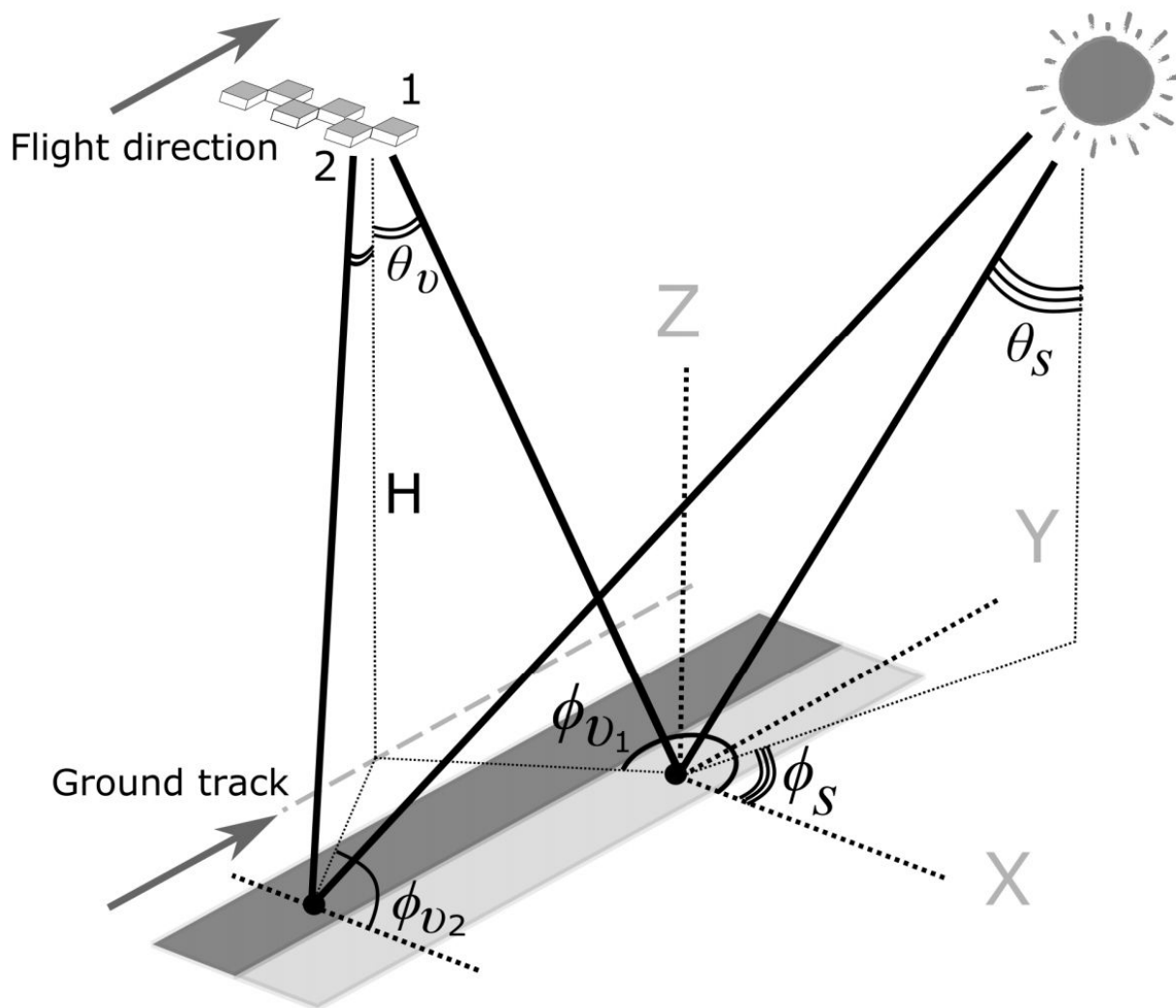


Optical waves remote sensing with Sentinel-2 (ESA-Copernicus) and Landsat-8 (NASA-USGS)

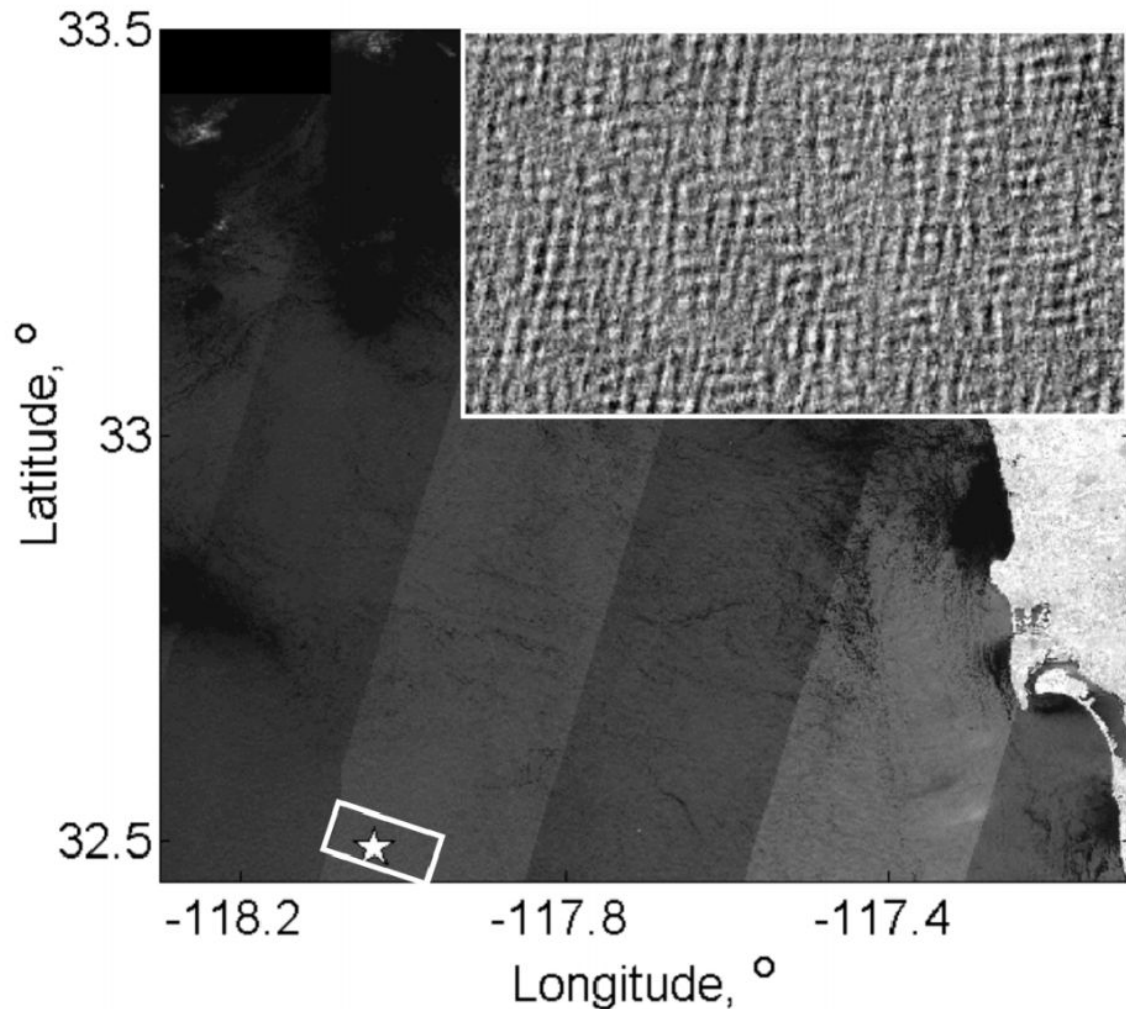




Sentinel-2



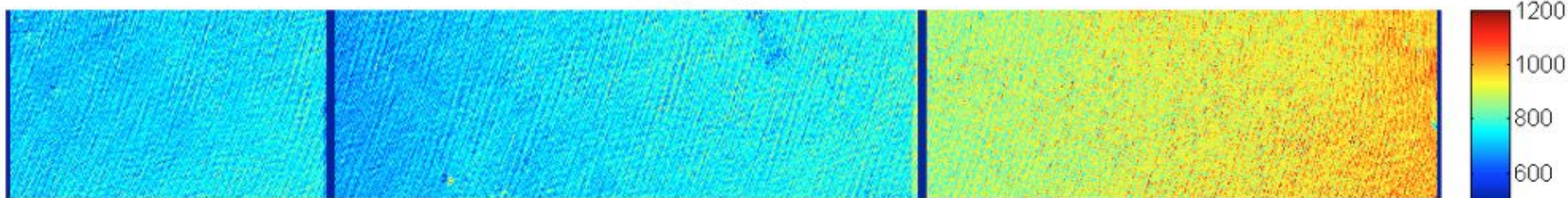
Sentinel-2



Sentinel-2 MSI image channel B04 off the California coast, 2016-04-19 627 18:44. The white star indicates location of the buoy 46086 - San Clemente Basin, 628 National Data Buoy Center (NDBC).

From brightness to slope

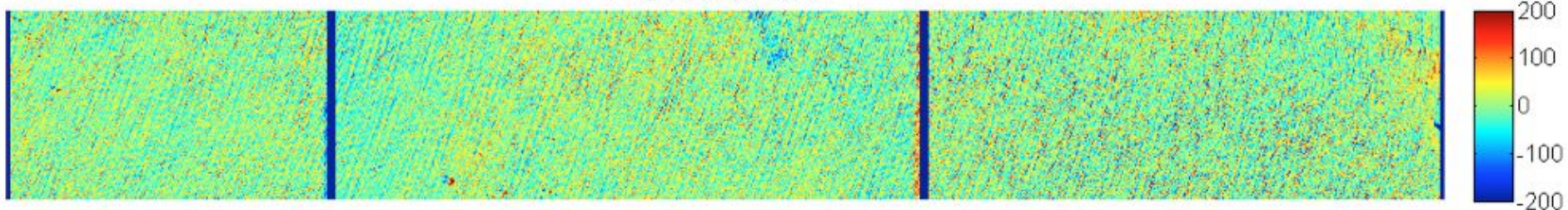
Image Fragment



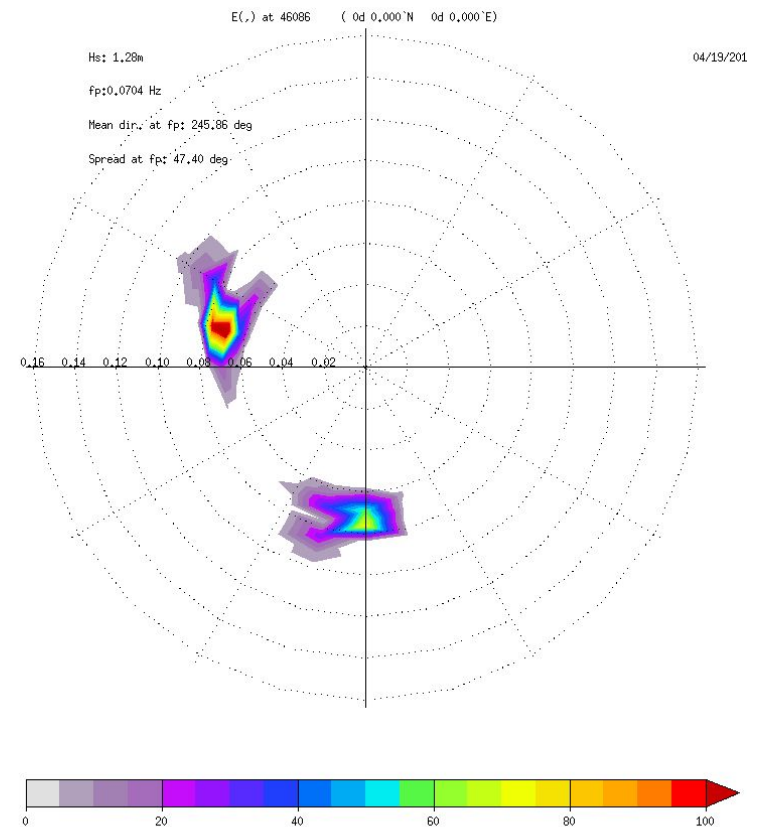
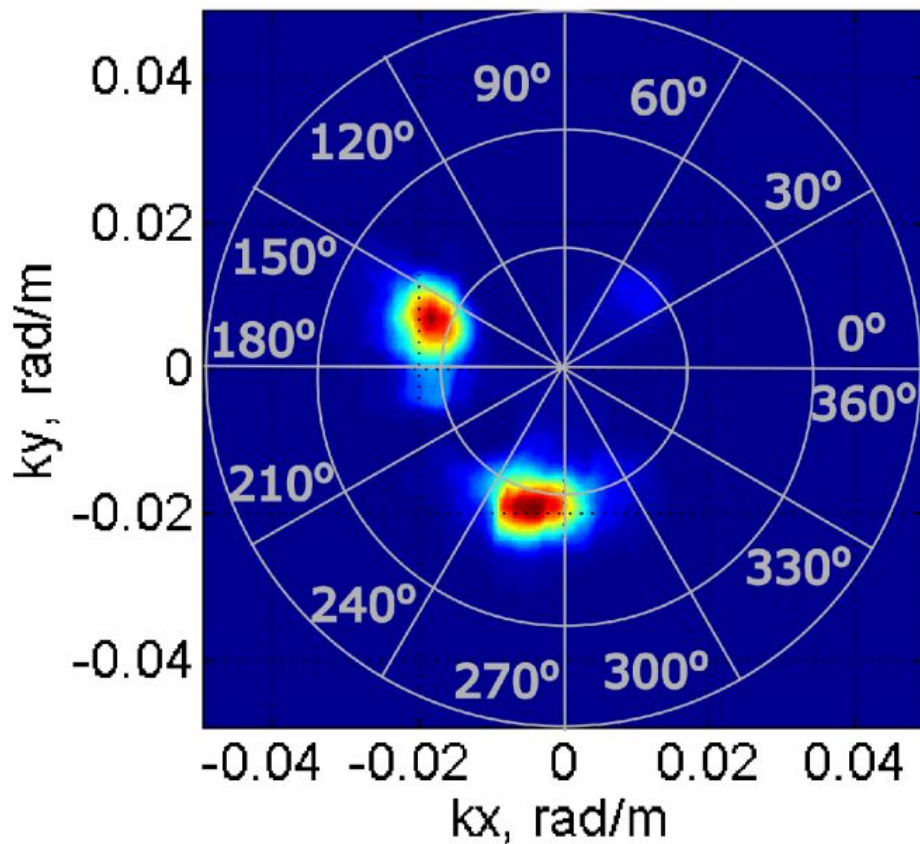
Mean Brightness



Brightness Variations



Comparison with buoy



Kudryavtsev, V., Yurovskaya, M., Chapron, B., Collard, F. and Donlon, C. (2017), Sun glitter imagery of ocean surface waves: 1. Directional spectrum retrieval and validation. J. Geophys. Res. Oceans. Accepted Author Manuscript. doi:10.1002/2016JC012425